EXERCISE 5

AIR PRESSURE AND WIND

THEME:

Air pressure is a major cause of the winds and is responsible for fair and inclement weather.

TERMS:

Air pressure Weather station Air temperature Wind direction Wind speed Anticyclone Wind arrow

Fair weather Cyclone Low pressure centers High pressure centers Isobars Inclement weather

GUIDELINES: At the conclusion of this exercise, you should be able to:

- 1. Interpret a weather station symbol. 2. Identify high and low pressure cells.
- 3. Explain the appearance of the sky in terms of fair weather and inclement weather.

Everything seems to be under a lot of pressure these days-even the atmosphere. Fortunately, most people are not particularly responsive to changes in air pressure. However, even small AIR PRESSURE changes can result in cloudiness and rain. A way to examine air pressure and its role in determining the weather is to analyze a map which depicts air pressure and winds.

The place to start an analysis is at the WEATHER STATION where the readings are taken. The station is represented by a circle. The weather data at the station is placed around the circle which is known as the station symbol (Figure 5.1). Pay particular attention to the units of measurement of each weather element.

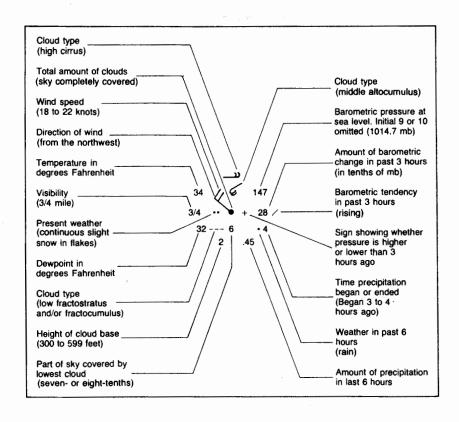


FIGURE 5.1 Weather station symbol.

An abbreviated station symbol would indicate current AIR TEMPERATURE, AIR PRESSURE, WIND DIRECTION, and WIND SPEED (Figure 5.2).

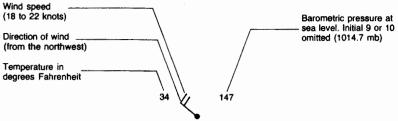


FIGURE 5.2 An abbreviated station symbol

The location of the data around the station symbol is important. By placing each of the values in a specified place, one can quickly locate the desired information.

Of particular interest is the representation of wind information. WIND DIRECTION is depicted as a straight line which is known as a WIND ARROW. The arrow flies with the wind. Thus, using the station symbol as a 360° circle, we can accurately place the arrow. For example, in Figure 5.2 the wind is coming

from 315°.

WIND SPEED is depicted by placing feathers on the arrow shaft (Figure 5.3).

1/2 feather represents 5 knots

Full feather represents 10 knots

Pendant represents 50 knots

(1 knot = 1 nautical mile/hour or 1.15 statute miles/hour)
 (1 knot = 1.85 kilometers/hour or 0.5 meters/second)

FIGURE 5.3 Wind arrow feathers.

In order to provide a clear view of the air pressure conditions on the map. points of equal air pressure are connected with a line. According to the isoline concept studied in Exercise 2, this line is known as an ISOBAR. Given enough weather station information, isobars typically close and form an oval. By looking at the value of numbers in the ovals and comparing their value to adjacent ovals, we can identify HIGH PRESSURE CENTERS and LOW PRESSURE CENTERS (Figure 5.4). This means that high and low pressure centers are always identified as such in relation to each other.

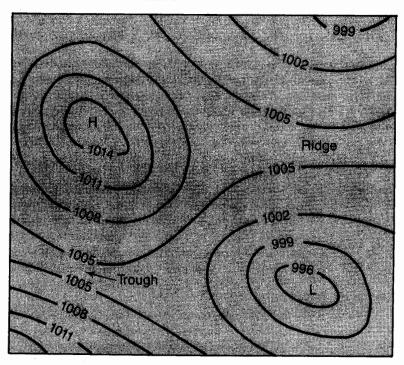


FIGURE 5.4 High and low pressure centers.

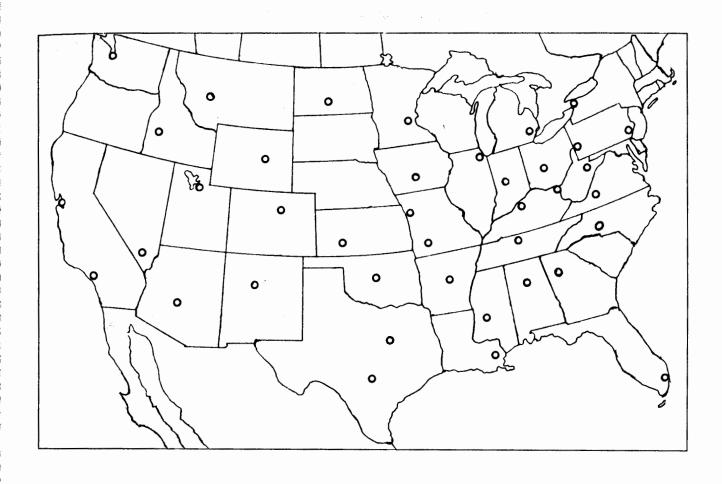
In the northern hemisphere, the air, e.g. wind, within a high pressure cell moves out from the center, down the pressure gradient, and clockwise. The cell or system is identified as an ANTICYCLONE. In the southern hemisphere, the motion of the air is counterclockwise. In either hemisphere, the anticyclone indicates FAIR WEATHER.

On the other hand, the air, or wind, of a low pressure cell in the northern hemisphere moves toward the center, lifts upwards as it converges and travels in a counterclockwise direction. It is known as a CYCLONE. The southern hemisphere air movement in a cyclone is clockwise. Whether northern or southern hemisphere, the cyclone means INCLEMENT WEATHER. As a result one can expect cloudy, if not stormy, weather.

The following activities will give you experience in working with weather data and weather map construction. Thus, when you look up at the sky you will be able to visualize a weather map or when looking at a weather map, you will know what the sky will look like once outdoors.

ACTIVITIES: AIR PRESSURE AND WINDS Class: Name:_____ 1. Construct a weather map using the data given below and the map of the United States. a. Plot the station symbols first. b. Next, draw isobars with a 4 millibar interval, e.g. 1000, 1004, 1008, etc. c. Place an "H" in the center of the high pressure cell. d. Place an "L" in the center of the low pressure cell. 2. What is the general trend of air movement around the high pressure center? 3. What evidence supports your observation about the air movement given in No. 2? 4. What is the general direction of air movement around the low pressure center? 5. What evidence supports your observation about the air movement given in No. 4? 6. Analyze the map, and indicate the weather of Cincinnati, Ohio in terms of: a. Wind direction _____ b. Wind speed _____ c. Pressure _____ d. Temperature _____ e. Is it fair or inclement? _____ f. Sky appearance _____

Location	Temp	Press (mb)	Wind direction	Wind Speed (knots)
Minneapolis, MN	43	1008	SW	10
Detroit, MI	45	1007	NE	10
Des Moines, IO	39	1012	W	10
Chicago, IL	4.5	1007	NE	10
Indianapolis, IN	48	999	NE	10
Columbus, OH	46	1003	E	10
Roanoke, VA	50	1008	SE	10
Springfield, MO	46	1012	NW	. 5
Kansas City, MO	48	1014	NW	5
Louisville, KY	57	997	NE	5
Nashville, TN	68	999	NE	10
Birmingham, AL	70	1008	SW	15
Atlanta, GA	66	1010	S	10
New Orleans, LA	73	1014	sw	5
Jackson, MS	73	1010	SW	10
Miami, FL	75	1017	SE	5
Charlotte, NC	66	1008	SW	15
Pittsburgh, PA	46	1008	NE	10
Philadelphia, PA	45	1016	SE	5
Buffalo, NY	43	1013	SE	5
Portland, ME	32	1018	0	Ō
Little Rock, AR	63	1011	SW	10
Fort Worth, TX	64	1014	NE	10
San Antonio, TX	79	1013	SE	5
Oklahoma City, OK	52	1015	NW	5
Dodge City, KS	43	1016	S	10
Bismarck, ND	46	1015	N	5
Helena, MT	57	1015	SW	15
Casper, WY	48	1019	NW	15
Boise, ID	55	1023	W	10
Salt Lake City, UT	52	1023	SE	10
Denver, CO	45	1019	N	5
Albuquerque, NM	57	1018	NW	10
Phoenix, AZ	61	1018	E	5
Las Vegas, NV	59	1022	E	10
Los Angeles, CA	61	1022	S	5
San Francisco, CA	57	1025	NW	5
Seattle, WA	52	1029	SW	5



EXERCISE ON ATMOSPHERIC STABILITY AND CLOUD DEVELOPMENT

At the end of this exercise, the student should be able to:

- 1. identify the components of the adiabatic process;
- 2. determine air mass stability and instability;
- 3. recognize the differences between the various adiabatic rates; and
- 4. calculate the lifting level of condensation to locate the base of a cloud.

Introduction

Weather conditions are determined to a great extent by the conditions of *stability* in the atmosphere. One afternoon in summer may be fair because stable air is present; another may have a thunderstorm because the air is unstable. In unstable conditions, pilots may warn their passengers that their passage over the unstable area may result in a bumpy ride, as plumes of air, called thermals, rise and strike the bottom of the aircraft.

Meteorologists rely on instruments, such as *radiosondes*, to measure weather elements above Earth's surface. From data on temperature at different altitudes, they can calculate the extent of atmospheric stability. Cloud types also offer a reliable means to identify atmospheric zones of instability. When the tops of fair weather cumulus clouds do not expand very far above their flat bases, the air is known to be stable. But if the tops of the clouds rise in great dome-shaped masses turbulent air will be located at great heights, and a storm is likely.

A meteorologist's predictions of visibility, ceilings, cloud cover, and storm conditions are based on information about the degree of stability of the atmosphere. The purpose of the exercise is to acquaint the student with the methods used to determine atmospheric stability.

The primary source for heating the atmosphere is the surface of Earth. As a result, the vertical temperature gradient for stationary air (no vertical movement) usually decreases with elevation. Atmospheric soundings (from radiosondes) over land and water throughout the year and from around the world indicate an average decrease in temperature of 3.5° F for each 1000 feet (0.65° C per 100 m) of altitude. This average value is known as the average temperature lapse rate or the normal lapse rate. Essentially, temperatures can vary throughout a stratified layer of air because of the addition or loss of heat energy. Daily soundings of the atmosphere reveal, however, that the lapse rate for stationary air varies considerably from the average. The values derived from real time measurements of the atmosphere are called the observed temperature lapse rate. Temperature readings obtained for each level of the atmosphere yield the environmental temperature lapse rate.

For the following problems, all temperatures are in Fahrenheit units, and all altitudes are in feet.

Part I—Normal Lapse Rate Calculations. Using the normal lapse rate, determine the following:

- (1) What is the temperature at an altitude of 10,000 feet if the sea level temperature is 60° F?_____.
- (2) Plot the temperature change with altitude on the first graph that follows this exercise. adiabatic paperate
- (3) The following temperatures were observed at a weather station during the early morning and mid-afternoon.

ALTITUDE (ft)	TEMPERATURES (* F)				
	Early Morning	Mid-Afternoon			
11000	22	30			
10000	25	32			
9000	28	35			
8000	32	40			
7000	38	45			
6000	40	51			
5000	42	55			
4000	63	61			
3000	55	64			
2000	48	69			
1000	43	72			
Surface	45	75			

- (4) Plot the preceding temperatures as dots on the second graph that follows this exercise. Use a solid line to connect the dots to show the early morning lapse rate curve and a dashed line to show the mid-afternoon curve.
- (5) Explain the early morning temperature curve. What is responsible for the different environmental lapse rates?
- (6) The lapse rate is determined by finding the difference in temperature between two altitudes, such as between the surface (75° F) and 11,000 feet (30° F) and then dividing that value by the number of 1000-foot units (11). What is the observed temperature lapse rate for the mid-afternoon?
- (7) The following temperatures in degrees Fahrenheit are radiosonde data collected at a mid-latitude weather station in the Northern Hemisphere. The soundings were taken on three different days.

ALTITUDE (f	t)	TEMPERA'	TURE (° F)
:	Day 1	Day 2	Day 3
11000	19	-10	16
10000	25	-5	19
9000	31	0	24
8000	37	9	29
7000	43	10	32
6000	49	13	35
5000	55	16	41
4000	61	19	50
3000	67	21	48
2000	73	25	46
1000	79	28	52
Surface 8	5	32	56

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(8) Plot the preceding temperatures on the third graph that follows this exercise and draw the vertical temperature gradients. Use a solid line for Day 1, a dashed line for Day 2, and a dotted-dashed line for Day 3. Based on the data and temperature curves, determine the season of year for each of the data sets. Day 1: ______ Day 2: ______ Day 3: ______
Label each curve on the graph with the appropriate season.

Part II—Adiabatic Lapse Rates. In Part I, you calculated the normal lapse rate, which results from temperature changes in stratified air by the addition or subtraction of heat energy. In Part II, you will calculate the adiabatic rates which are associated with ascending and descending masses of air and the responses of the air masses to changes in pressure. In nature, a volume of air may be forced upward into the atmosphere in one of three ways: frontal lifting, convectional lifting, and orographic lifting. As a mass of air ascends, it expands as atmospheric pressure decreases and becomes less dense and therefore, cooler (cooling by expansion). A descending mass of air will undergo compression beneath the increasing pressures nearer to the surface (heating by compression).

The rates of temperature change depend on phase changes of water in the air. If air ascends but has not cooled to its dew point (it is not saturated), it will cool at a constant rate of 5.5° F for each 1000 feet of ascent (or 1° C per 100 m). This rate is known as the dry adiabatic lapse rate. If the mass of air should cool to its dew point, condensation will occur, releasing heat energy (the latent heat of vaporization). With the release of heat energy, the rising air will continue to cool, but at a lesser rate, only 3.2° F for each 1000 feet of ascent (or 0.58° C per 100 m). This diminished cooling rate is called the moist (or wet) adiabatic lapse rate.

The level at which condensation begins (the lifting level of condensation, or LLC) is at the altitude where the temperature and dew point temperature are equal. The rate at which the dew point drops is 1° F for each 1000 feet of elevation (or 0.18° C for each 100 m). The lifting level of condensation approach rate can be determined by the formula:

Temperature - Dew Point × 222.22 feet per ° F = LLC

(1) Calculate the LCL for the following. You will be given the surface temperature and the surface dew point. Under "Cloud Composition" write liquid or solid (ice).

Site	Surface Temperature (° F)	Surface Dew Point (° F)	LCL (ft)	Cloud Composition
a)	68°	50°		
b)	90°	40°		
c)	40°	38°		

- (2) On the chart below, assume a surface temperature of 80° F and a dew point of 42.2° F, and calculate the following:
 - a) Calculate the normal lapse rate of 6.5° per 1000 feet.
 - b) Calculate the LCL and draw a line across the graph to indicate it.
 - c) Calculate the dry adiabatic rate of 5.5° F per 1000 feet up to the LCL, and calculate the wet adiabatic lapse rate of 3° F per 1000 feet from the LCL to 10,000 feet.
 - d) Calculate the difference between the normal lapse rate and the adiabatic rate (wet and dry) for every 1000 feet. Place these figures on the line provided. Use a plus sign (+) when the parcel of air has a higher temperature than the surrounding (ambient) air temperature (normal lapse rate) and use a minus sign (-) when it is lower.
 - e) What is the condition of stability shown? _____.

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