EXAMPLE ARTICLE Summary R. Akers 3-10-83 Summary

Baldwin, E.M., 1974, Eocene stratigraphy of southwestern Oregon: Oregon Dept. of Geology and Mineral Industries, Bulletin 83, 40 p.

Baldwin gives a detailed description of the Eccene formations in south-western Oregon and uses the revised stratigraphy in order to reconstruct the paleo-geography of the Eccene coastal margin. Klamath pre-Tertiary strata occupy only a small part of the study area (Fig. 1); however, Baldwin concludes that they were a major source area for sediments and their Cenozoic tectonic history is closely related to that of the Eccene sedimentary basins of western Oregon.

The Eccene formations are dominantly sandstones and siltstones with some conglomeratic and coaly beds (Fig. 3). The earliest Eccene formation, the Rose-burg Formation, was closely folded and thrust eastwards soon after deposition. This early Eccene telescoping of Roseburg strata may have been in response to subduction of an oceanic plate to the west. The Paleocene to early Eccene seaway (Fig. 6) in which the Roseburg Formation was deposited reached northward into Washington and volcanic and sedimentary rocks of the Crescent Formation may be correlative.

Deformation of the Roseburg Formation resulted in a brief erosional event prior to the deposition of the overlying Lookingglass Formation. The Lookingglass Formation unconformably overlies the Roseburg Formation and onlaps upon the pre-Tertiary strata along the periphery of the basin, toward the Klamath Mountains (Fig. 6). A brief period of erosion (probably due to crustal uplift), after the deposition of Lookingglass strata, created an unconformity upon which the Flournoy Formation onlapped during the middle Eocene (Fig. 2). The Flournoy seaway was one of the most restricted during the Eocene (Fig. 6). Conglomerate, pebbly sandstone, and coal were deposited in shallow seas with interfingering non-marine strata. The finer grained, thin beds of the upper part of the Flournoy Formation imply deposition in a quiet, deeper neritic environment. The source for the Flournoy strata is questionable; however, some evidence indicates that they may have come from the Klamath Mnts. as end-filling a north trending basin.

The middle Eocene Tyee Formation rests unconformably on the Flournoy, Lookingglass, and Roseburg strata and shows a conspicuous lack of basal conglomerate; indicating that the source area was not adjacent to the basin and that the sediment was delivered by rivers bearing sand and silt. Previous workers (Snavely, 1964; Lovell, 1969) suggest that the sediments came from the south (Klamath province) and were carried northward over a submarine surface that was previously subdued by erosion (Flournoy-Tyee unconformity). The southern portion of this elongate basin (Fig. 13) is dominated by nonturbidite facies while the northern portion is dominated by turbidite deposits. A broad transitional zone between these two lithofacies is present. An offlap towards the north and west restricted the size of the basin toward the end of Tyee deposition and the beginning of Elkton deposition (Fig. 13).

The Elkton Formation is gradational above the Tyee Formation, becoming finer grained upward. The Elkton beds represent deposition at a time when the energy of the streams was diminishing. The overlying Bateman Formation represents the final deposits of the offlapping sea (Fig. 13). Uplift and erosion occurred prior to the deposition of the Coaledo and Spencer Formations. The Coaledo and Spencer Formations were deposited in shallow, encroaching seas during the late Eccene. Sediments for these formations came from the Klamath province and from a possible land barrier that existed at the time of deposition (Fig. 16). The non-marine, late Eccene Colestin Formation consists of lava flows, tuffs, and tuffaceous sandstones and conglomerates that outcrop along the western edge of the Cascade range. The Colestin appears to have been marginal to the late Eccene marine Coaledo strata, but the exact relationships are unknown. The Bastendorff Formation is predominantly shale, whose coarser, near-shore equivalents have been eroded away (Fig. 16).

Basin morphology, numerous unconformities, and northerly directed sediment transport suggest that sedimentation during the Eocene was controlled by tectonism. Although this view has yet to be proven.

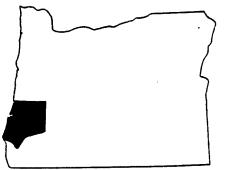
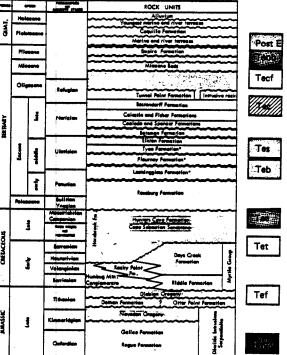


Figure 1. Location map of study area. Pre-Tertiary strata found in southern of study area.

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Figure 2. Strettgraphic chart for south



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Post Eocene

Bastendorff Formation:
Thinly bedded gray to buff shale and minor sands

Colestin and Fisher Formations:
Flows and volcaniciastic beds, co
tuffaceous sandstons

Coaledo Formation:

Sequence consisting of three members; upper members consist of coal-bearing, cro fuffaceous sandstone, and middle member thin-bedded sittstone with minor eandstone.

Spencer Formation: Sendstone with some co-

Bateman Formation:
Thickly bedded to misseesus, deltaic sen

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Elkton Formation:

IL FORMATION:
Indistinctly bedded micaesous siltstone with intermittent beds of messive sandstone.

Tyee Formation:
Thick sequence of mythmicelly bedded, micecous studence and siltatons; coal-bearing at Eden ridge.

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Flourney Formation:
Rhythmically bedded micaecous candstone po
upward into thin-bedded candstone and siltstone.

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Lookingglass Formation: Rhythmically bedded emdstone and siltstone; basal beds are coal-bearing and conglomeratic locally near the base of the section.

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Roseburg Formation:
Thick sequence of sandstone and siltatone; rhythmically bedded locally; contains minor conglomerate and measure sandstone; pillowed and breceited submarine baselts (Terv) are abundant locally.

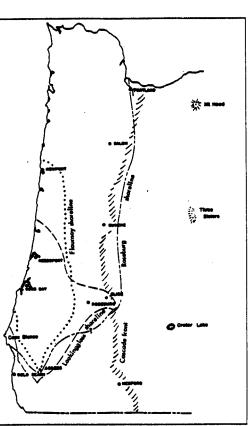
Tertiary intrusive rock:

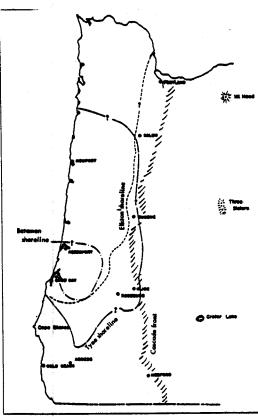
Sills and dikes of basic to intermediate composition.

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Pretertiary

Figure 3. Description of Eccene strata in southwestern Oregon.





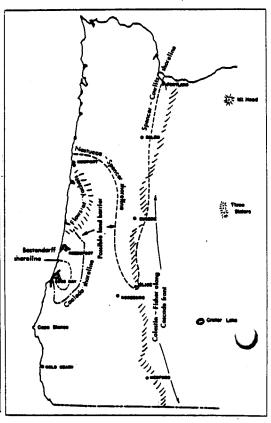


Figure 13. Paleageographic map of Tyee, Eikton, and Bat

Figure 16. Paleogeographic map of Coaledo, Colestin, Sp