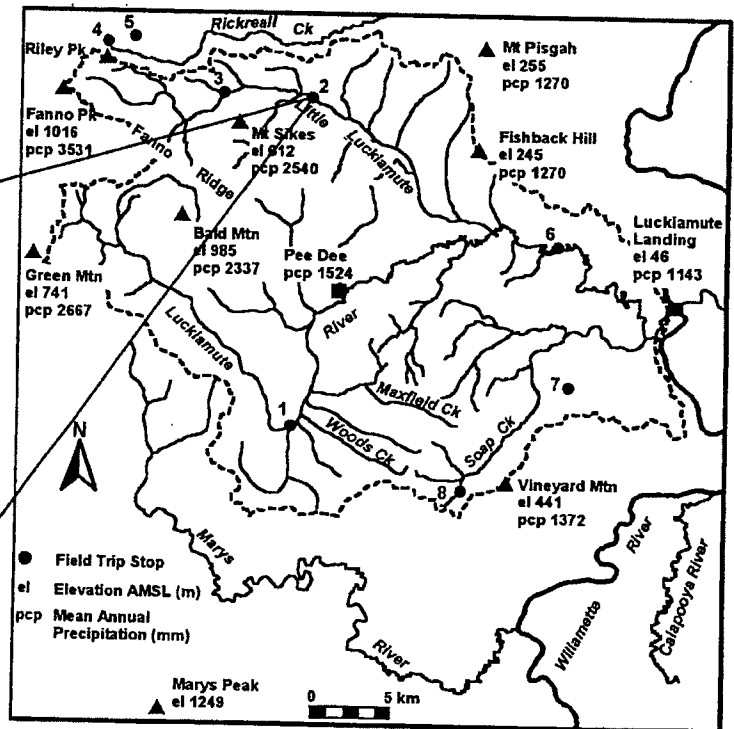


Field Guide to the Luckiamute River Watershed, Upper Willamette Basin



Trip Leader :

**Steve Taylor, Ph.D., Earth and Physical
Science Dept., Western Oregon University**

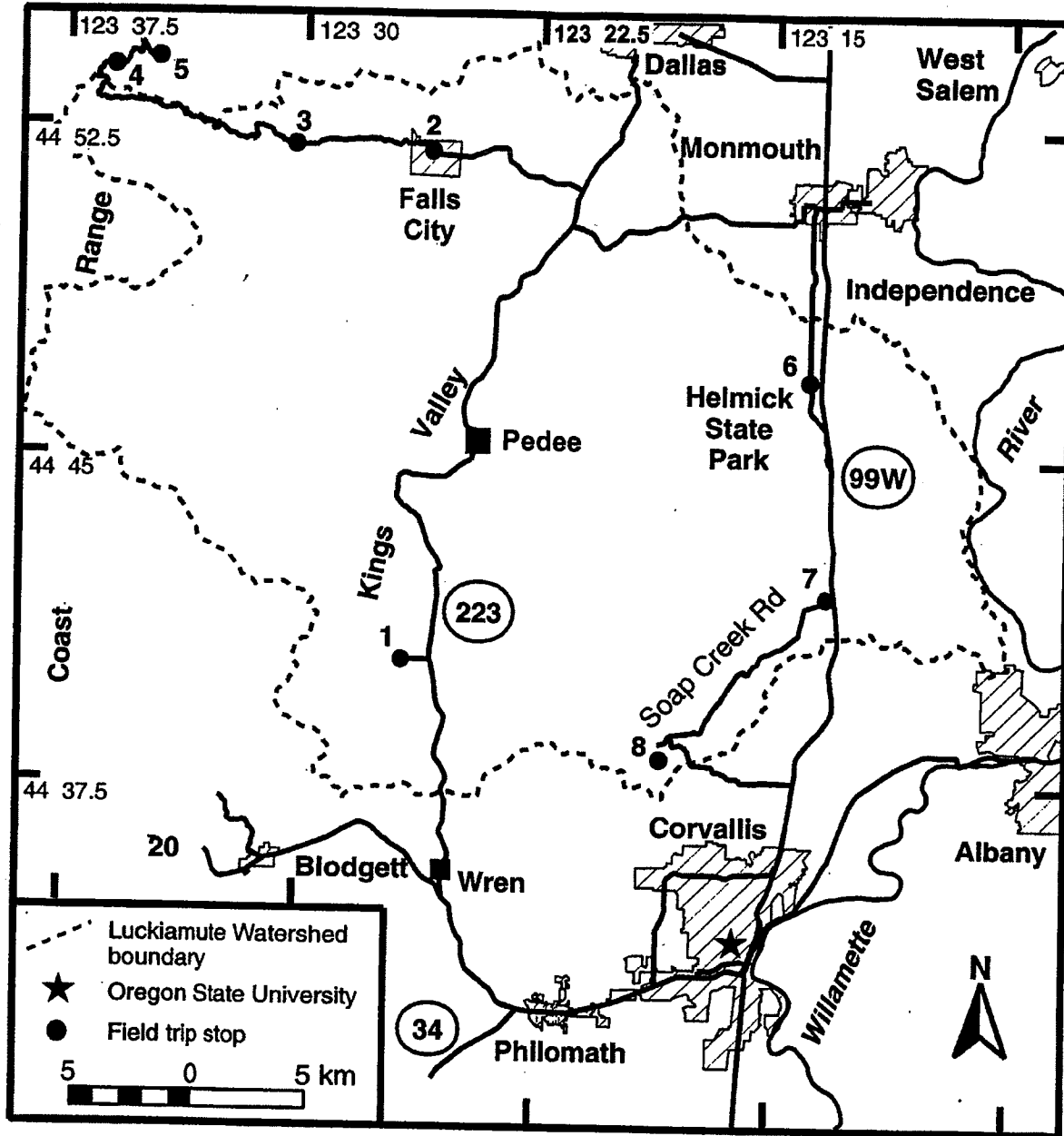


Figure 1. Location map and field trip route for the Luckiamute Watershed.

River to 1016 m (3333 ft) at Fanno Peak. The Luckiamute has an average gradient of 3 m/km, a total stream length of 90.7 km, and an average basin elevation of 277 m (910 ft) (Rhea, 1993; Slack and others, 1993). Fanno Ridge separates the watershed into two tributary subbasins, with the Little Luckiamute to the north and the main stem of the Luckiamute to the south (Kings Valley) (Fig. 2). Lower-order tributaries include Boughey Creek, Waymire

Creek, Vincent Creek, Plunkett Creek, Woods Creek, Maxfield Creek, and Soap Creek.

The greater Willamette Valley extends northward 190 km from Eugene to Portland, Oregon. This lowland is up to 60 km wide, separating the Coast Range to the west from the Cascade Range to the east. Valley floor elevations range from 150 m (500 ft) to 3 m (10 ft), with an average gradient of 2 m/km (Slack and others, 1993).

Physiographic Setting of the Luckiamute Watershed

•Boundaries

- Crest of Coast Range to West (headlands)
- Willamette River to East

•Drainage Area = 815 km²

- Largest Fifth-Field Watershed in central and northern Coast Range

•Primary Tributaries

- Little Luckiamute – northern watershed
- Luckiamute – southern watershed

•Secondary Tributaries

- Soap Creek, Maxfield Creek, Woods Creek, Teal Ck

•Elevation Range:

- Min: 46 m (150 ft) at Willamette
- Max: 1016 m (3333 ft) at Fanno Peak
- Avg. Basin Elevation: 277 m (910 ft)

•Basin Morphometry

- Average Stream Gradient: 3 m /km
- Total Stream Length: 90.7 km

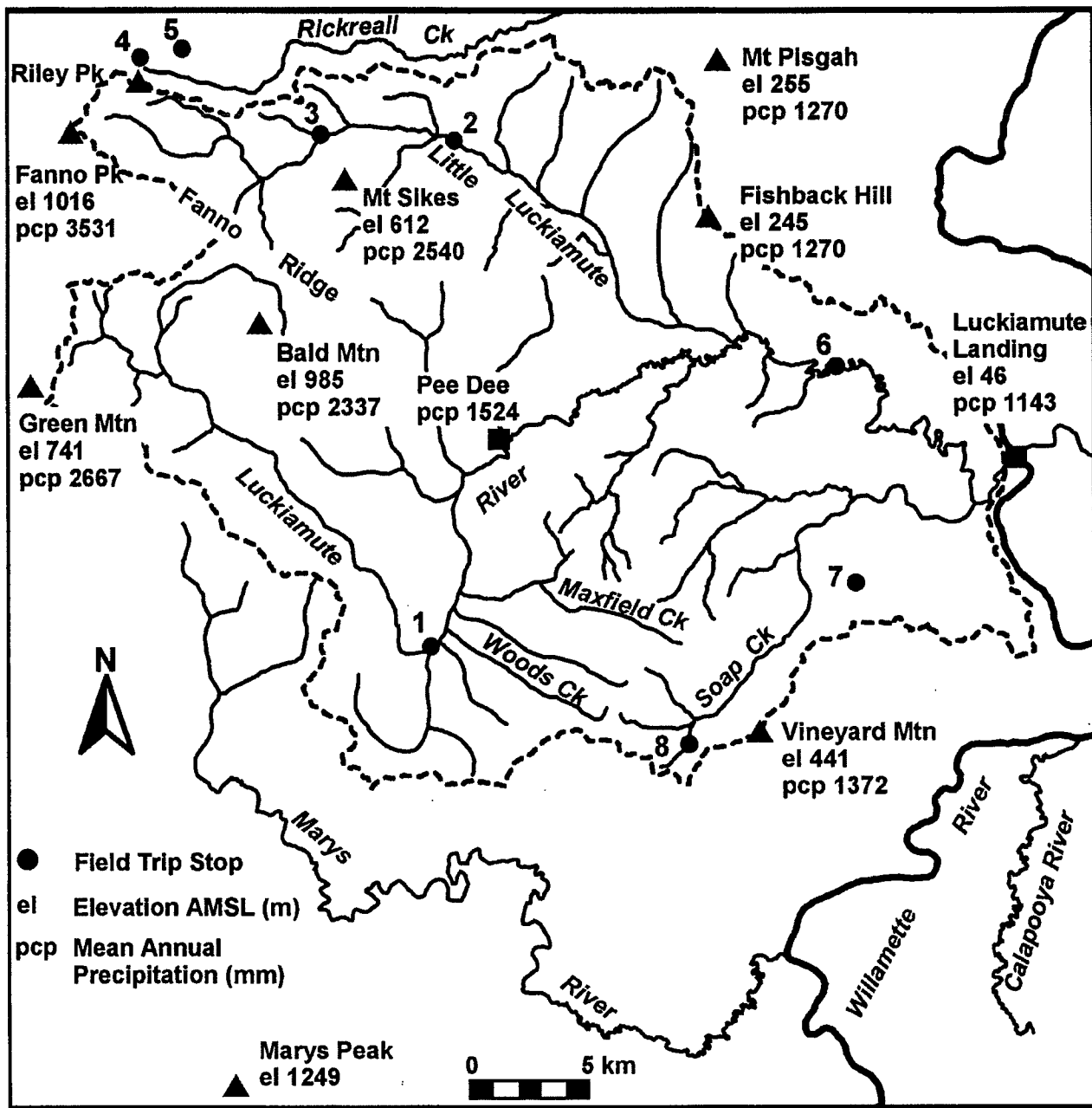


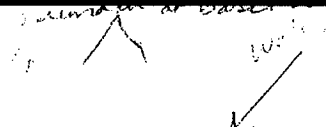
Figure 2. Physiographic map and spot annual precipitation for the Luckiamute Watershed.

Handwritten notes:
 1. C. Chubb
 YACM/5/0/010

Tectonic Setting of the Luckiamute Watershed

- Convergent Tectonic Margin
 - Subduction of Juan de Fuca Plate Beneath North America
 - Convergent Rates: 3.5-4.0 cm/yr
 - Style of Tectonism
 - Oblique Convergence
 - Tectonic accretion
 - Clockwise Rotation
- Coast Range Orogenesis
 - Accreted Marine Volcanic and Sedimentary Rocks
 - Active Uplift Between 15-10 Ma to Present
 - Neotectonics
 - General Uplift and Eastward Tilting
- Tectonic Influence on Luckiamute
 - Luckiamute drains the eastward tilted flanks of the Coast Range (Rhea, 1993)
 - Luckiamute Watershed located at segment boundary of Juan de Fuca Subduction zone

with 1000' elevation at base of Coast Range
Watershed



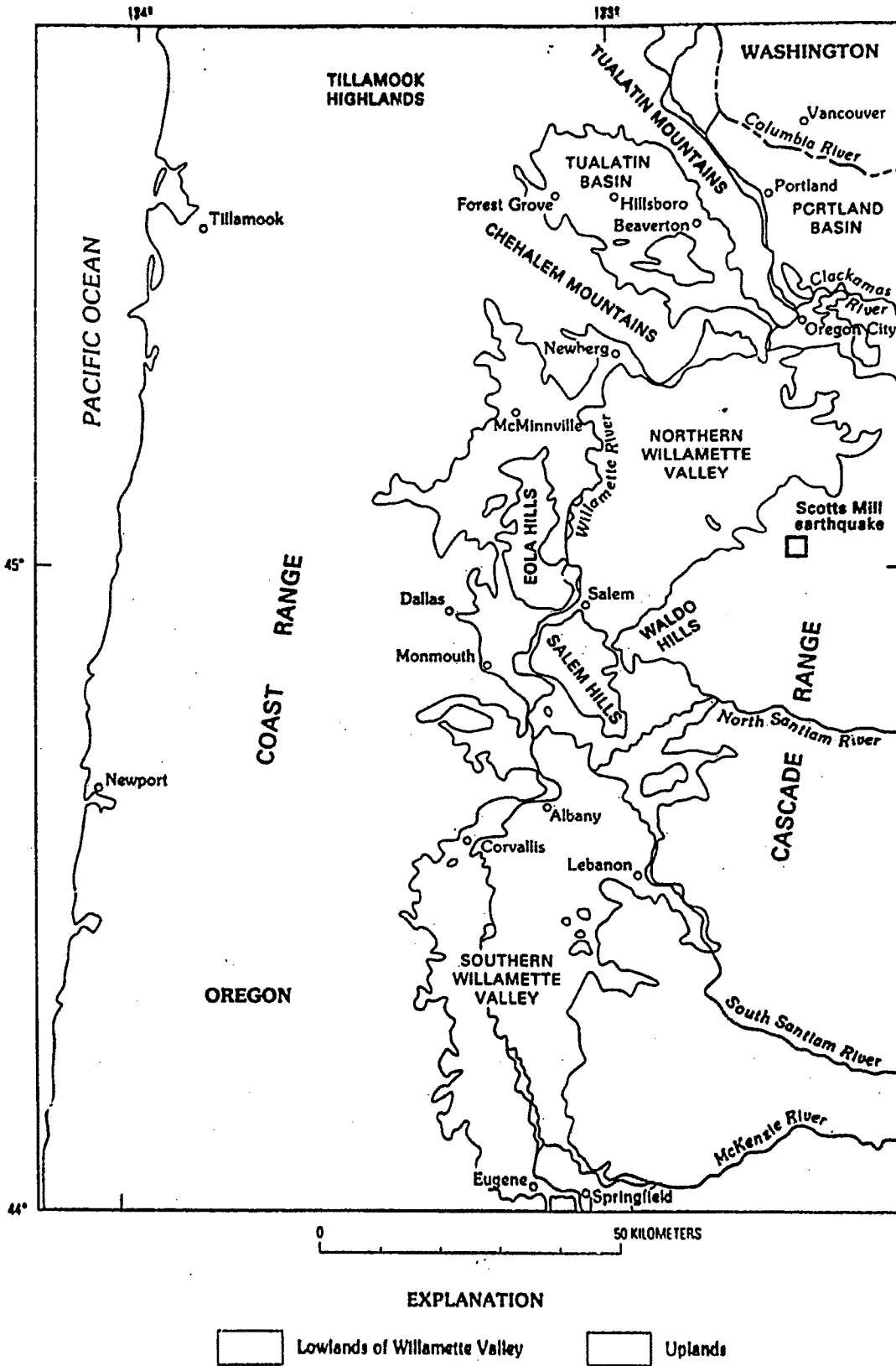
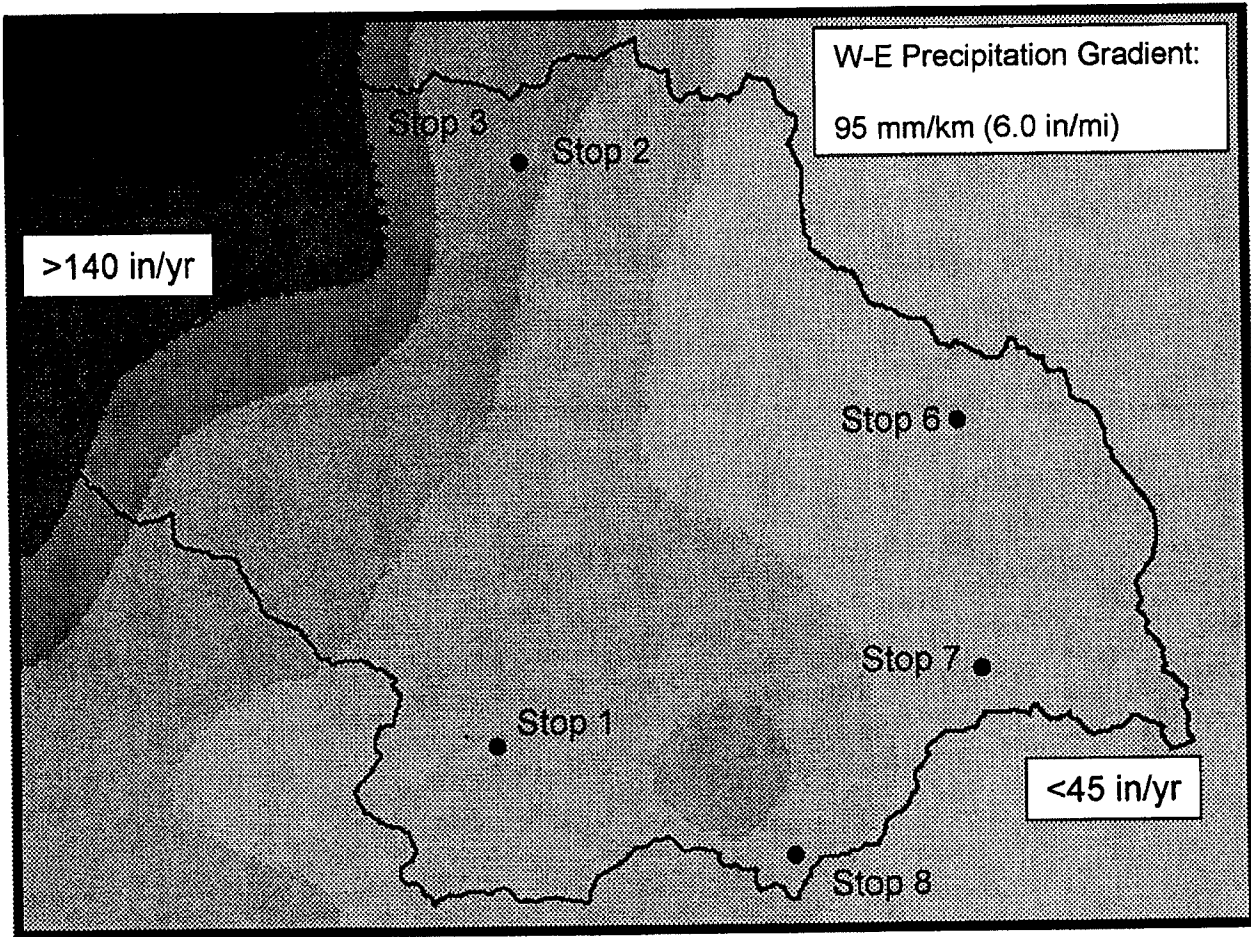


Figure 77. Geographic and physiographic features of the Willamette Valley and Portland and Tualatin basins, northwestern Oregon. The square indicates the epicenter of the March 25, 1993, Scotts Mills earthquake.

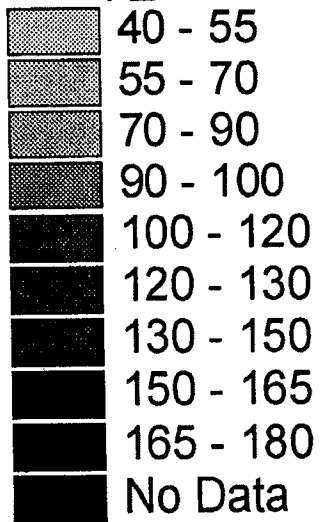
8



● Tripstop.shp

□ Luckbound.shp

Precip_90 (inches)

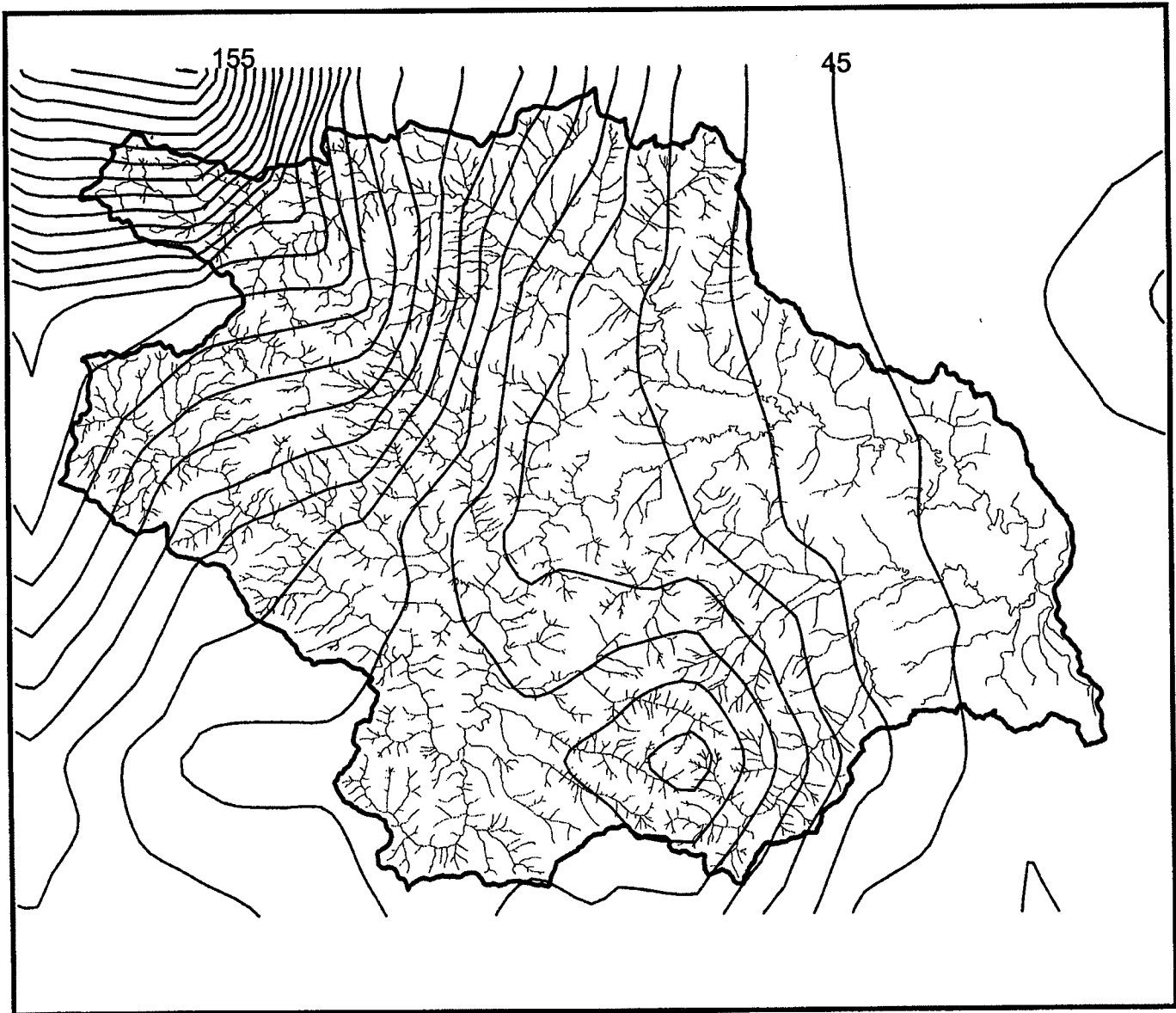


1990 Average Annual Precipitation for the Luckiamute Watershed (inches) (from the Oregon Climate Service)

Annual Precipitation

Basin Maximum: 3600 mm (>140 in) – Divide
 Basin Minimum: 1140 mm (~45 in) – Willamette Valley
 Basin-wide Precipitation Average = 1894 mm (~75 in)
 Seasonal Precipitation Cycle (October – March)

Luckiamute Basin Avg. Annual Rainfall (CI = 5 inches /yr)



5000 0 5000 10000 Meters

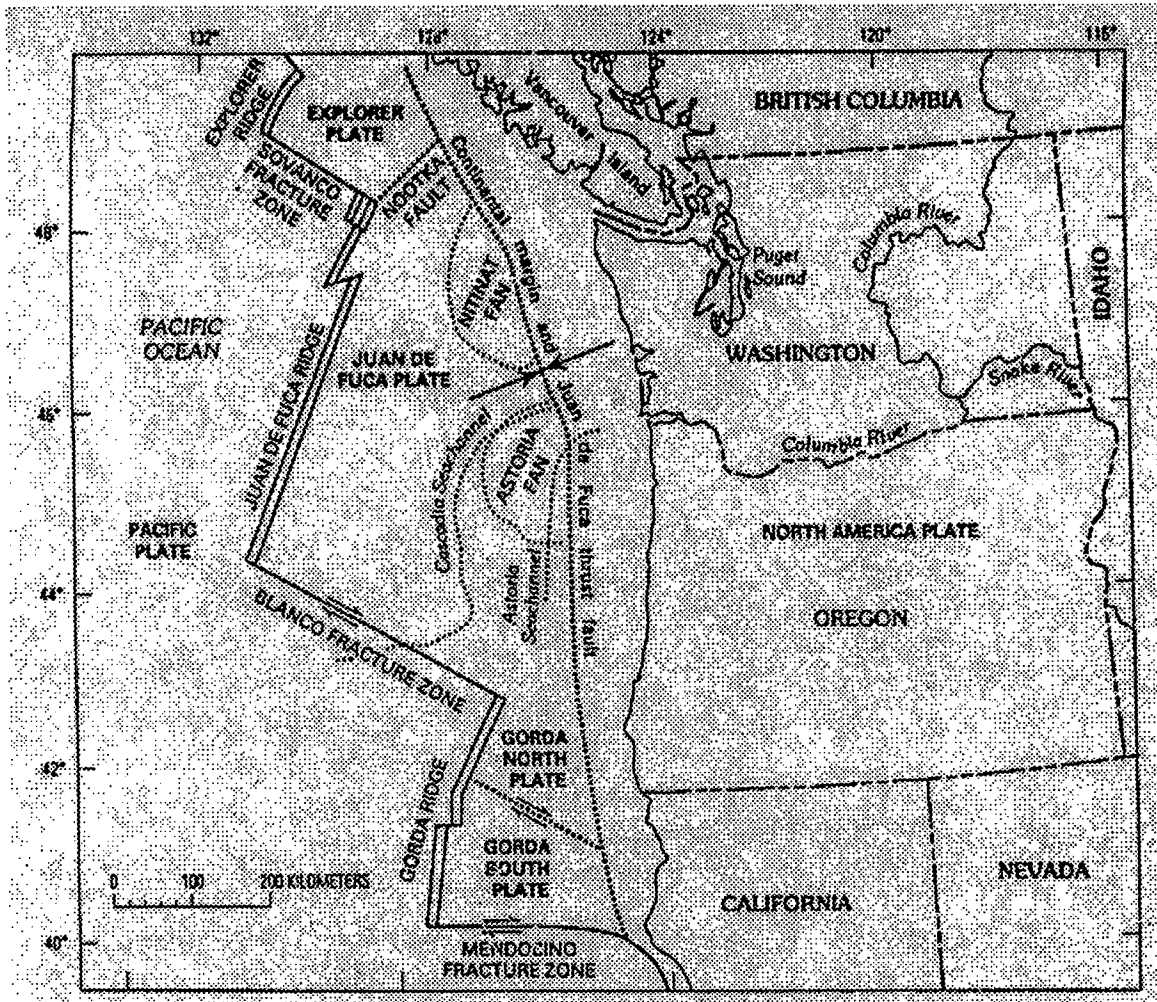
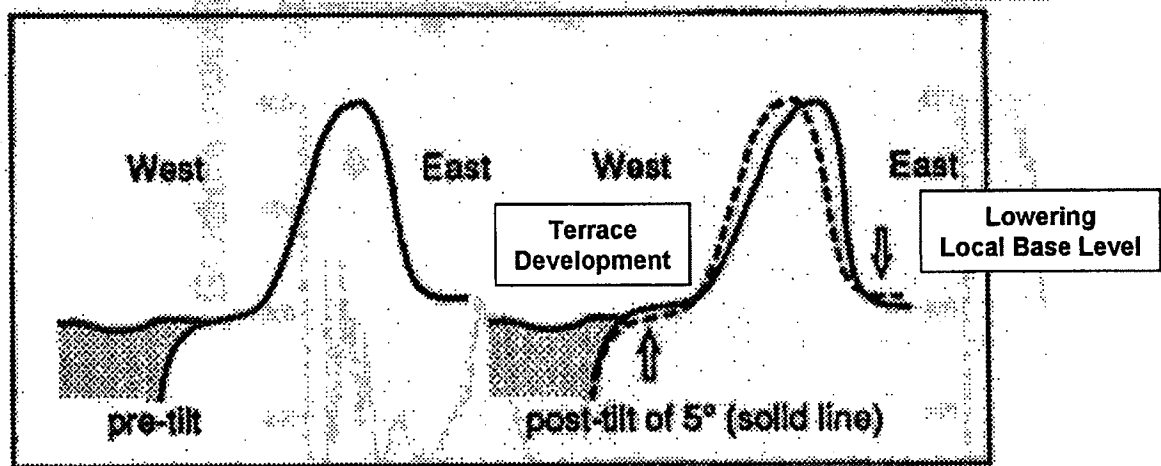


Plate tectonic configuration of the Pacific Northwest.



Cartoon showing effects of Coast Range tilting on watershed gradient (from Rhea, 1993)

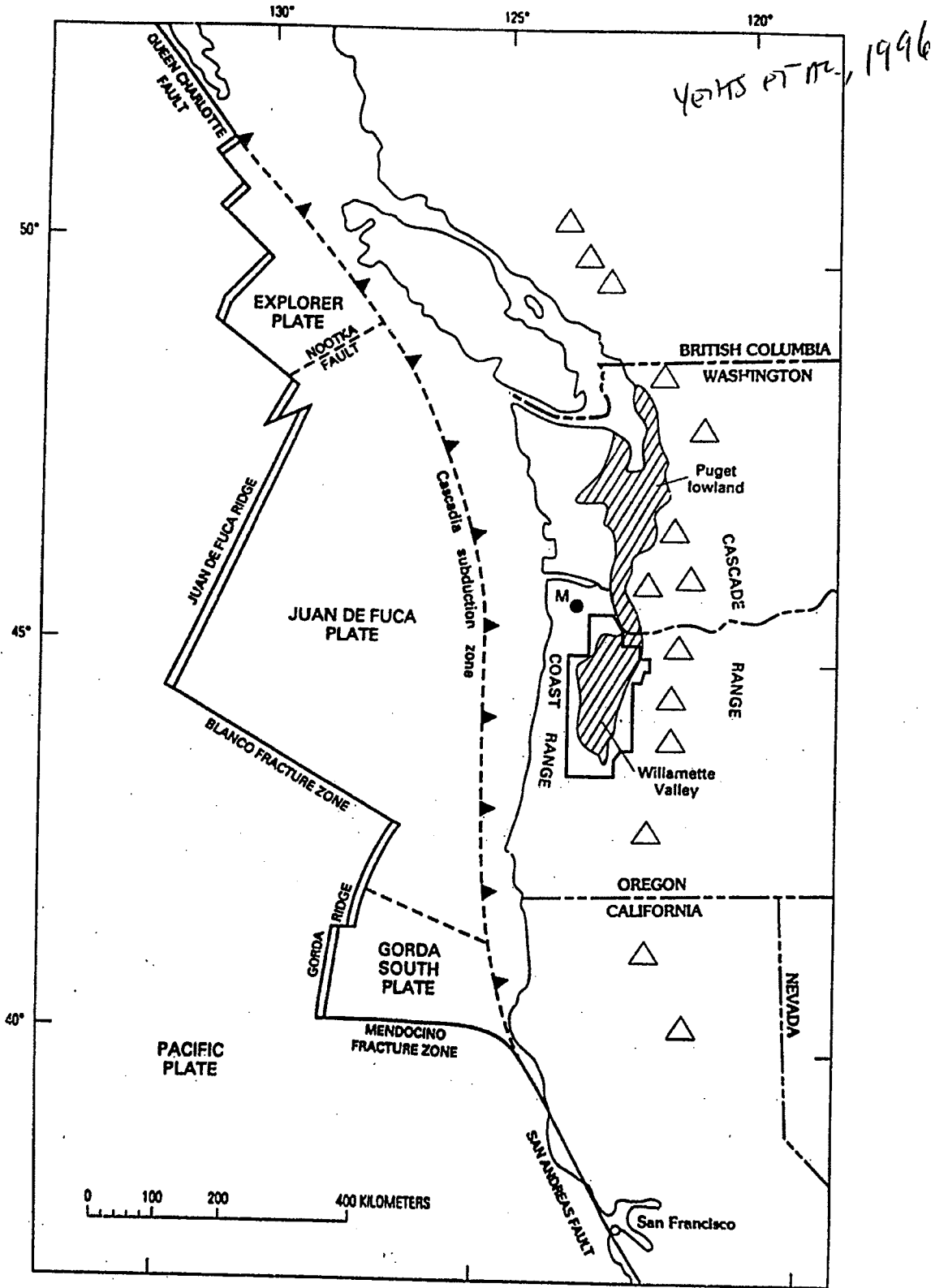


Figure 76. Plate boundaries of the Pacific Northwest showing locations of tectonic features and the Willamette Valley study area. Heavy line, study-area boundary; hatched area, Willamette Valley and Puget lowland; sawteeth denote upper plate of thrust fault. Major stratovolcanoes are shown by open triangles. Dot labeled "M" in northwestern Oregon is the Mist gas field.

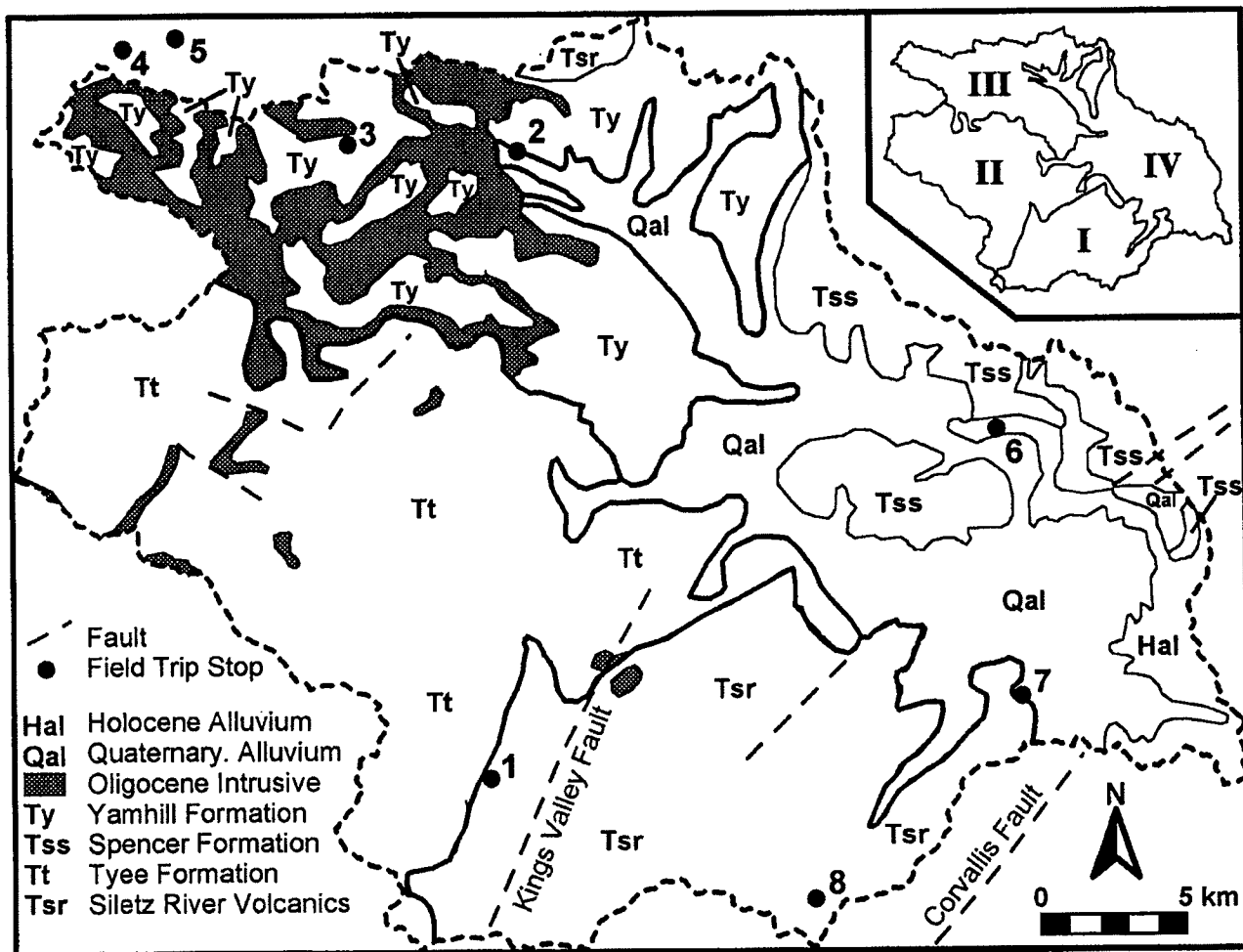
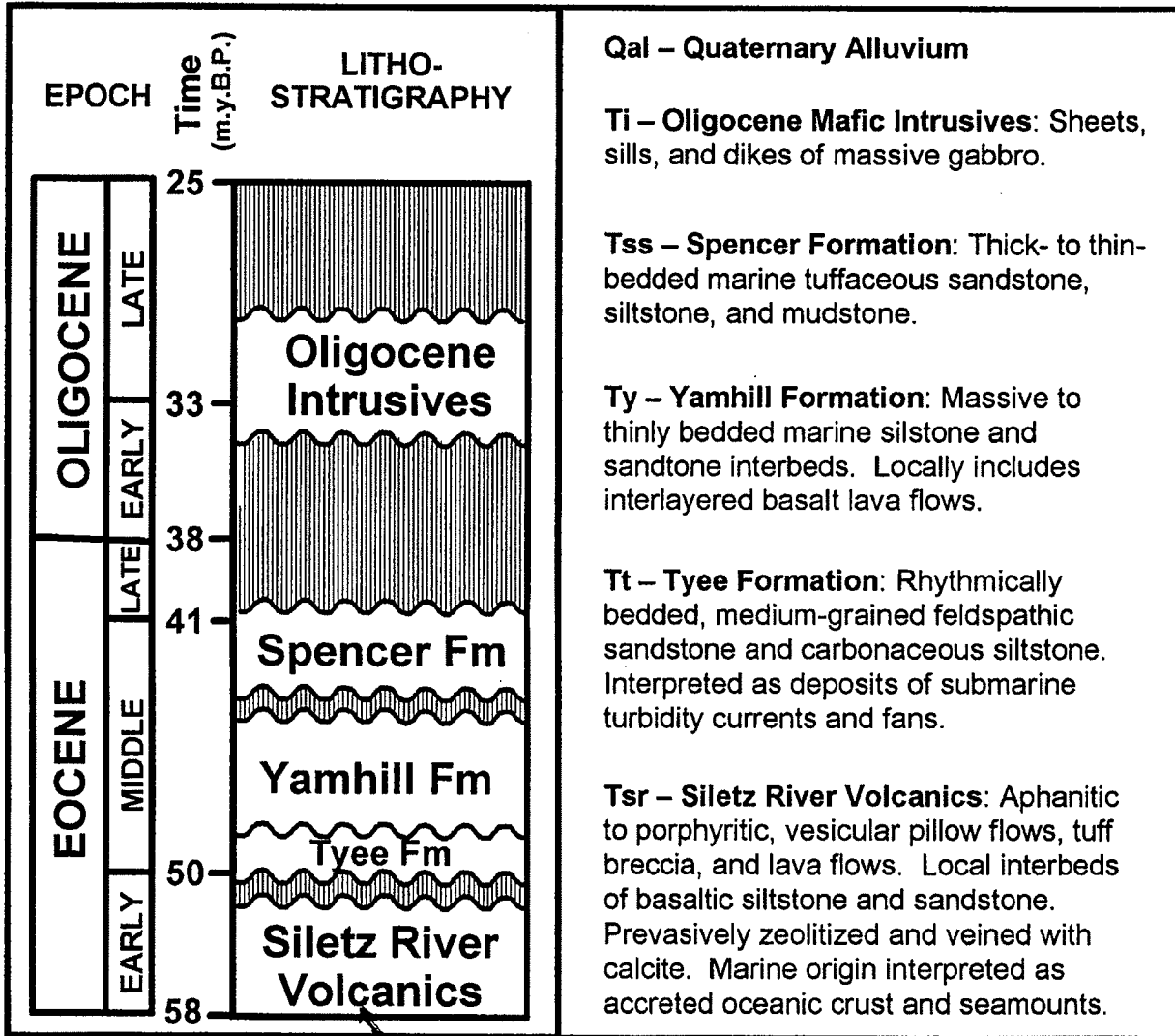


Figure 3. Bedrock geology of the Luckiamute Watershed (after Walker and MacLeod, 1991). Inset map shows grouping of recognized lithospatial domains: I = Siletz River Domain, II = Tye Domain, III = Yamhill-Ti (Tertiary Intrusive) Domain, IV = Spencer-Valley Fill Domain.

Bedrock Geology of the Luckiamute Watershed



Spencer rolling hills

step hills by road!!!

w/ both @ same elevation

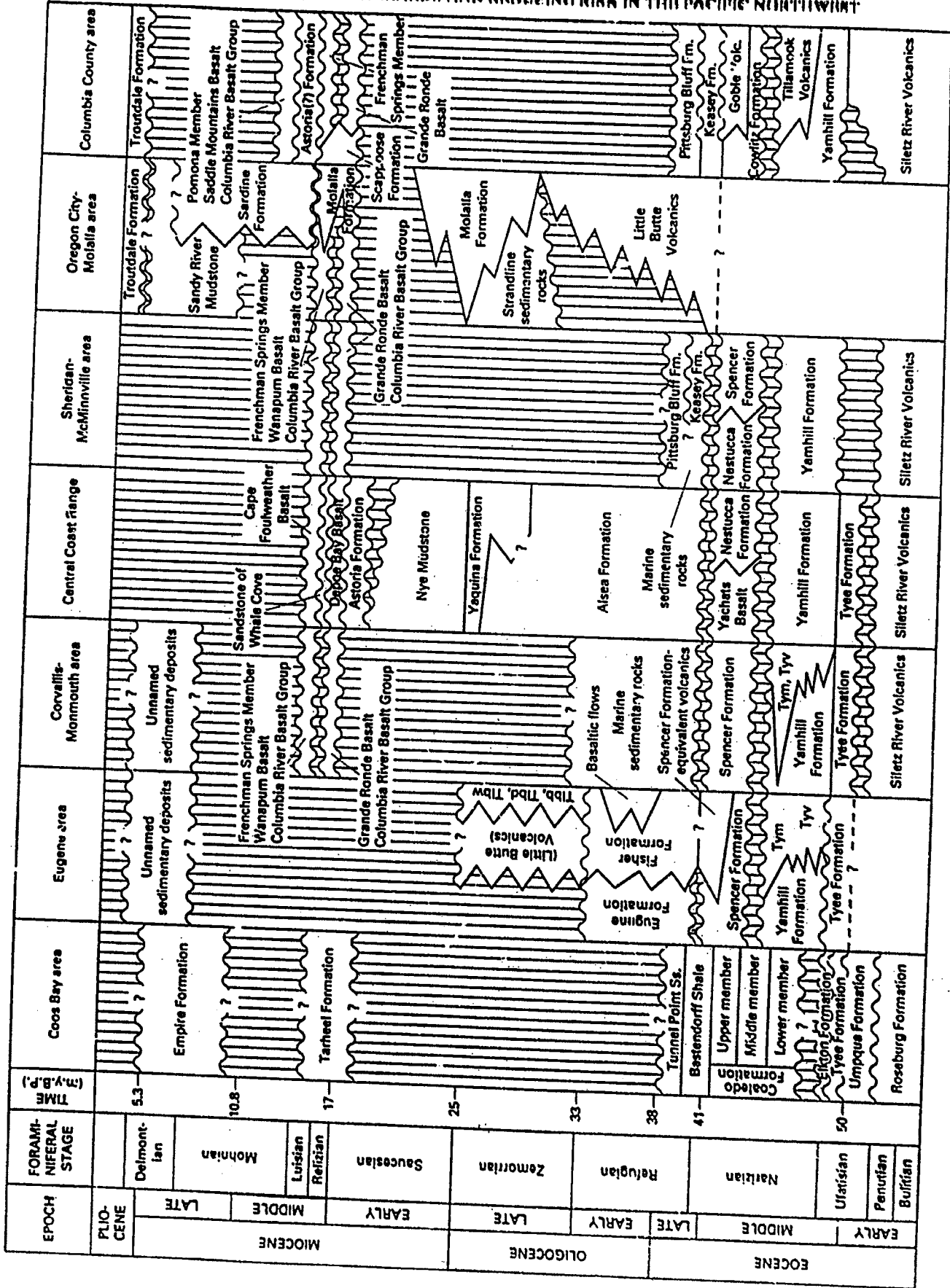
means for H. distance

uplift known

Flat floor plain siltstone - Maudslayi - 1000000000 - 10000000000

Maudslayi 1000

ASSESSING EARTHQUAKE HAZARDS AND BUILDING RISK IN THE PACIFIC NORTHWEST



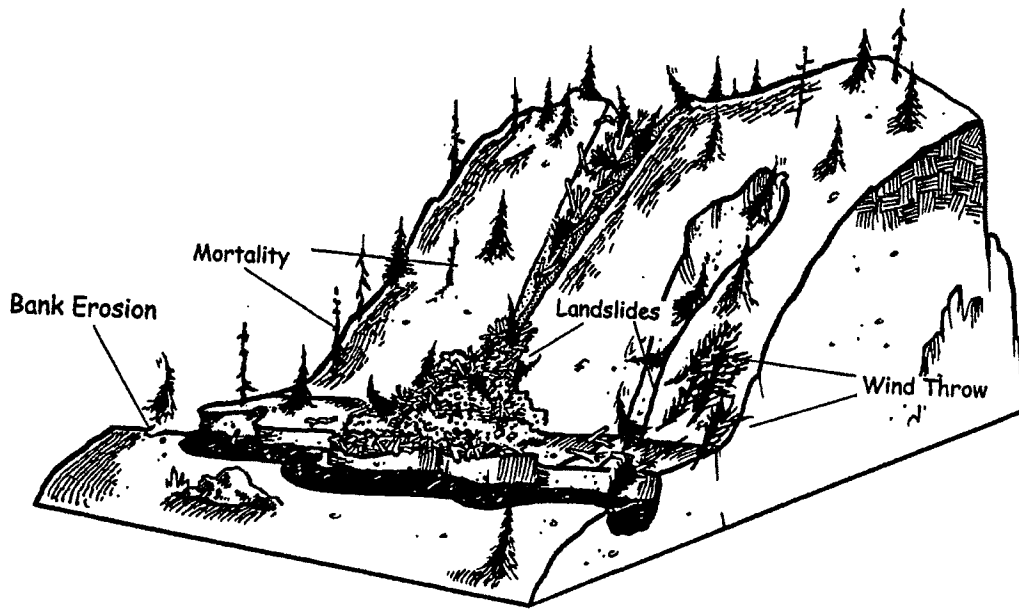
Geomorphology of the Luckiamute Watershed

Valley Floor-Fluvial Regime

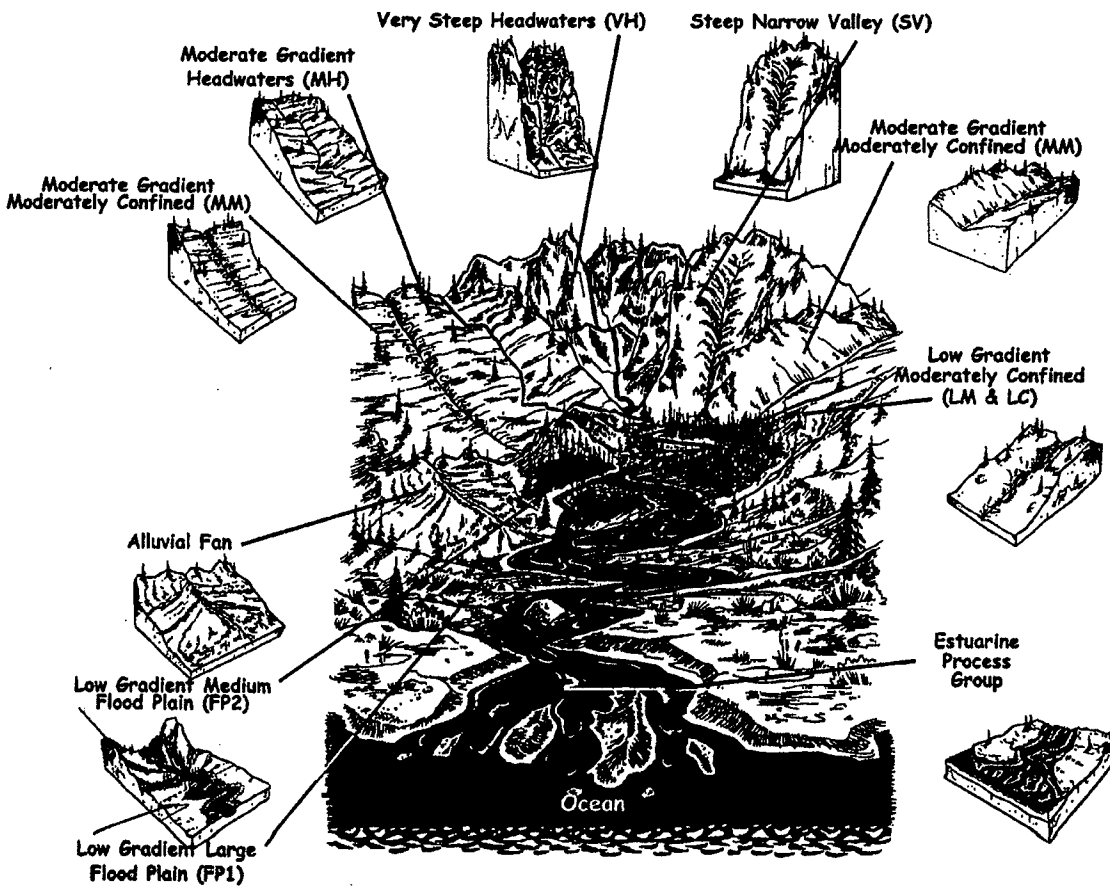
- Landforms
 - Channel
 - Floodplain
 - Terrace
 - Small-scale Fans
 - Strath-pediment surfaces
 - Low-relief colluvial hillslopes (Spencer Fm)
- Deposits
 - Alluvial Fill (sorted sand and gravel)
 - Debris Flow Deposits (diamicton)
 - Slackwater Silts and Clay
- Processes
 - Channel Transport
 - Overbank Sedimentation

Hillslope-Colluvial Regime

- Landforms
 - Ridge tops
 - Hillslopes-Sideslopes
 - Hollows
 - Pediment Surfaces
- Deposits
 - Colluvium (gravel diamicton)
 - Residuum (gravel diamicton)
- Processes
 - Colluvial Creep
 - Debris Slide / Flow
 - Tree-throw / Bioturbation



Geomorphic configuration of mountainous watersheds in the Coast Range.



Typical distribution of CHTs in a mountainous watershed.

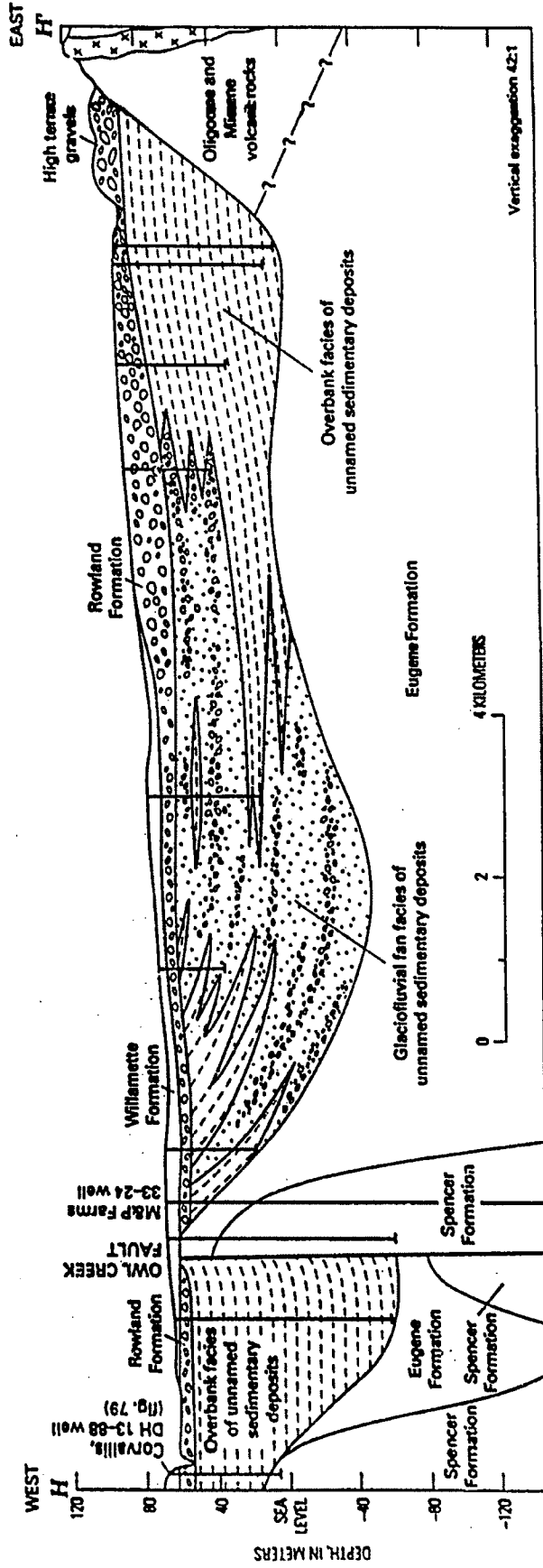


Figure 82. Structural cross section between Corvallis and Lebanon, Oreg., showing channel and overbank facies of unnamed fluvial sedimentary deposits, high-terrace gravels, late Pleistocene outwash deposits of the Rowland Formation, and catastrophic flood deposits of the Willamette Formation. Data are from water wells, engineering bore holes, and petroleum-exploration wells.

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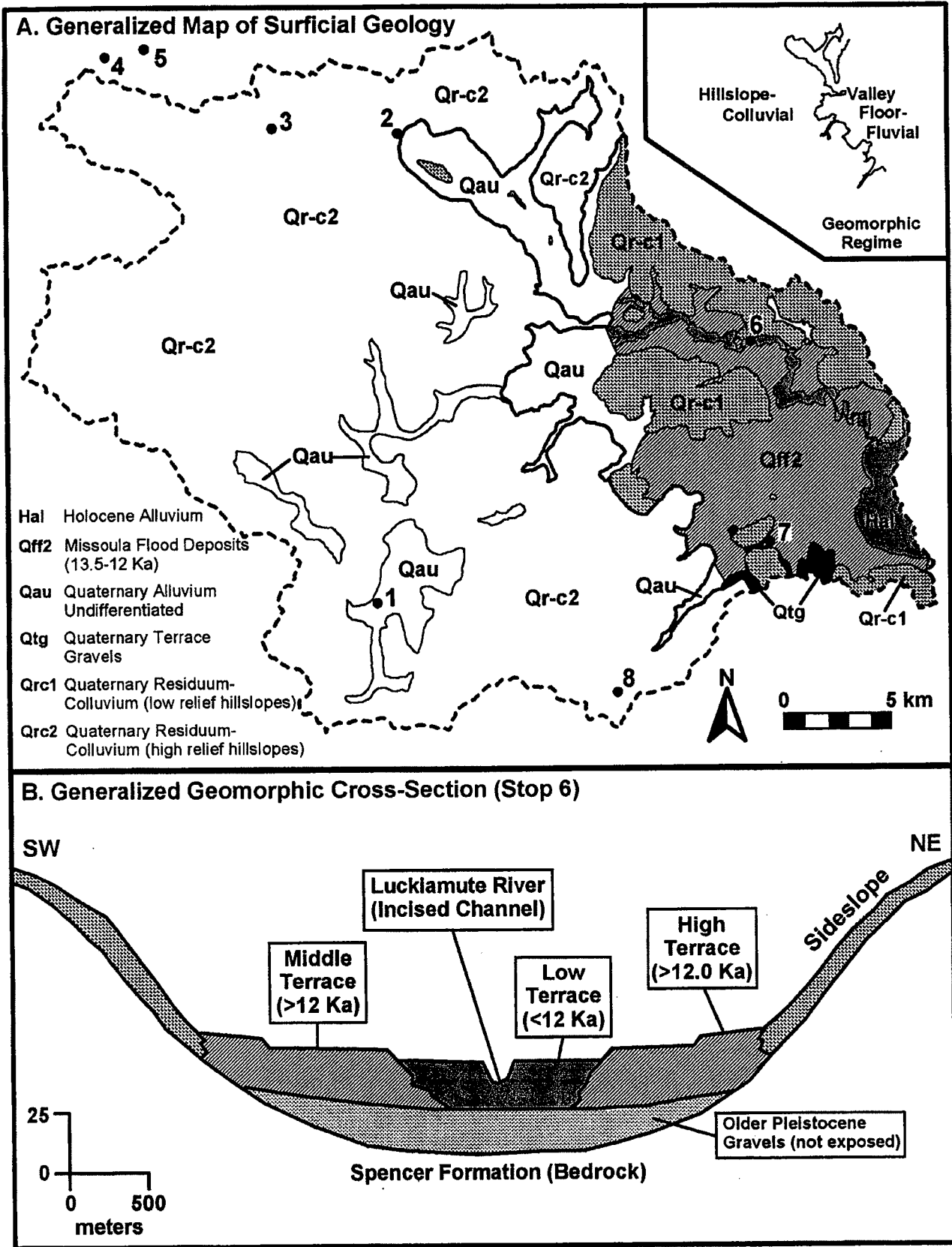
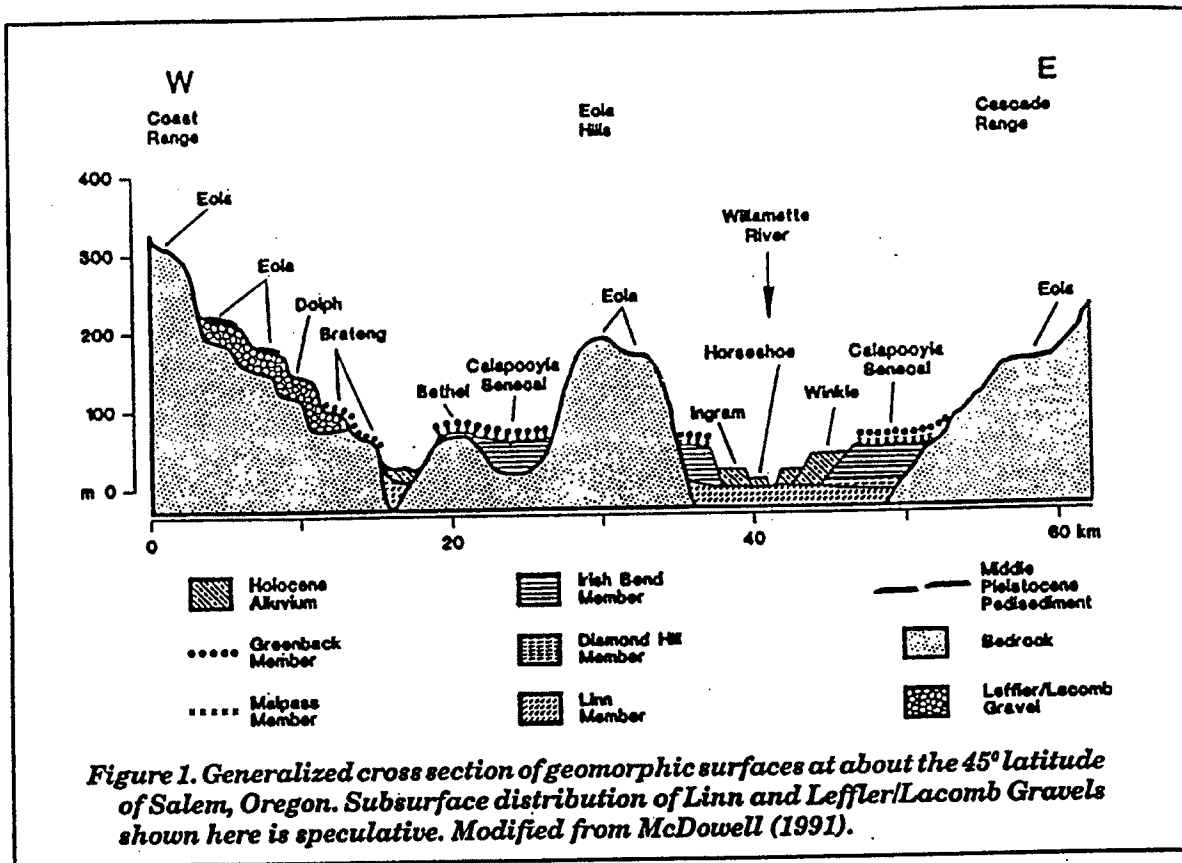


Figure 4. Surficial geology and geomorphology of the Luckiamute River Basin. Surficial map units are modified from O'Connor and others (2001), after Taylor and others (1996). Cross section shown in frame B represents generalized landform elements at Helmick State Park (Stop 6).



associated Greenback Member as described by Glassman and Kling (1980). Allison's (1935) original work reported no erratics lower than 30 m in the main valley but some as low as 10.5 m in the Portland area. Dozens of erratics locations are shown below 76 m and only rarely are erratics described at the maximum elevation of 122 m.

According to Allen et al. (1986), one small group of six boulders on Judkins Point east of Eugene occurs at an elevation of 198 m. It is assumed that these particular erratics are on the Eola surfaces rather than on Dolph Surfaces. Allen et al. (1986), attribute these high erratics to Indian transport, but an old Lake Missoula flood origin is also a possibility. Taken collectively, the data shows that erratics occur at various elevations above 30 m in the main valley, and they could be represent several different glacial Lake Missoula floods, rather than just the highest flood.

In general, the erratics range in size from small pebbles to boulders. They are composed of granites, granodiorites, gneisses, quartzites, schist, argillites, and phyllites. Allison (1935, 1936) noted at least 300

occurrences of erratics in the Willamette Valley. According to Allen et al. (1986), probably fewer than 50 of the erratic boulders still exist today.

The erratic rocks all require some outside source. Since the Willamette Valley lacks late Pleistocene continental till and associated outwash, and areas to the north show the definite presence of unmodified (by glaciers) Pleistocene non-glacial soil, it seems unlikely that the erratics were brought in by glaciers from the north. Allison (1935) related the erratics and the Willamette Silts, to be discussed later, to surges of glacial meltwater. This happened when the ice dams at the mouth of the Clark Fork River in Montana failed on different occasions, releasing the flow of Glacial Lake Missoula, to create the "Spokane Floods," between about 15 ka and 12.8 ka years before present (Allen et al., 1986).

The "Spokane Flood" (Bretz, 1919) was later named the "Missoula Flood," when the Montana Lake source for the flood water was discovered. It is commonly published (Allen et al., 1986) that, between 15 ka and 12.8 ka years ago, there were as many as 40 "Missoula Floods", so the term "floods" is

TABLE 3. GEOMORPHIC SURFACES OF THE WILLAMETTE VALLEY*

Geomorphic Surface	Surface Type	Dominant Surface-forming Process	Final Surface-forming Event	Major Associated Deposits	Approximate Age (B.P.)	Typical Soils
Horseshoe	Depositional	Fluvial lateral accretion	Fluvial lateral accretion	Recent alluvium	<300	Fluventic Haploxeroll
Ingram	Depositional	Fluvial lateral accretion	Fluvial lateral accretion	Recent alluvium	<5,000	Cumulic Ultic Haploxeroll
Winkle	Depositional	Fluvial lateral accretion	Fluvial lateral accretion	Recent alluvium	12,000 to 5,000	Pachic Ultic Argixerol
Champoeg	Depositional	Phase 2 flood proximal facies deposition	Phase 2 flood proximal facies deposition	"Lacustrine deposits"†	13,000	Argiaquic Xeric Argialboll, Aquic Xerocherpt
Senecal	Erosional	Local fluvial incision of Calapooyia surface	Deposition of Greenback Member	Willamette Formation	13,000	Aquiltic Argixerol, Pachic Ultic Argixeroll
Calapooyia	Erosional/ depositional	Plantation of Irish Bend Member and older deposits	Deposition of Greenback Member	Willamette Formation	13,000	Typic Albaqualf, Argiaquic Xeric Argialboll
Quad§	Uplifted(?)	Uplift of Willamette Formation deposits	Deposition of Greenback Member	Willamette Formation	13,000	Pacific Ultic Argixeroll
Bethel**	Depositional/ draped on erosional	Willamette Formation deposits draped on low bedrock knolls	Deposition of Greenback Member	Willamette Formation	13,000	Aquiltic Argixeroll
Brateng‡	Depositional/ draped on erosional	Greenback Member deposits draped on hillslopes	Deposition of Greenback Member	Greenback Member over Tertiary bedrock	13,000	Ultic Haploxeralf
Dolph	Depositional/ some erosional	Fluvial/glaciofluvial aggradation	Deposition of Greenback Member	Lacomb and Lefler gravels, with Greenback Member in places	Mid-Pleistocene or younger	Ultic Haploxeroll, Ultic Haploxeralf
Eola	Erosional	Pedimentation	Deposition of loess, gravels, or pedis sediment	Unnamed pedis sediment, Lacomb and Lefler gravels, Helvetia Formation, Portland Hills Silt§§	Early Pleistocene	Xeric Haplohumult
Looney	Erosional	Mass movement	Ongoing mass movement and slope erosion	Colluvium	Time-transgressive	Typic Haplohumult, Typic Haplumbrept

*Based on Balster and Parsons (1968, 1969) and Parsons and others (1970), except as noted.

†Schlicker and Deacon (1967).

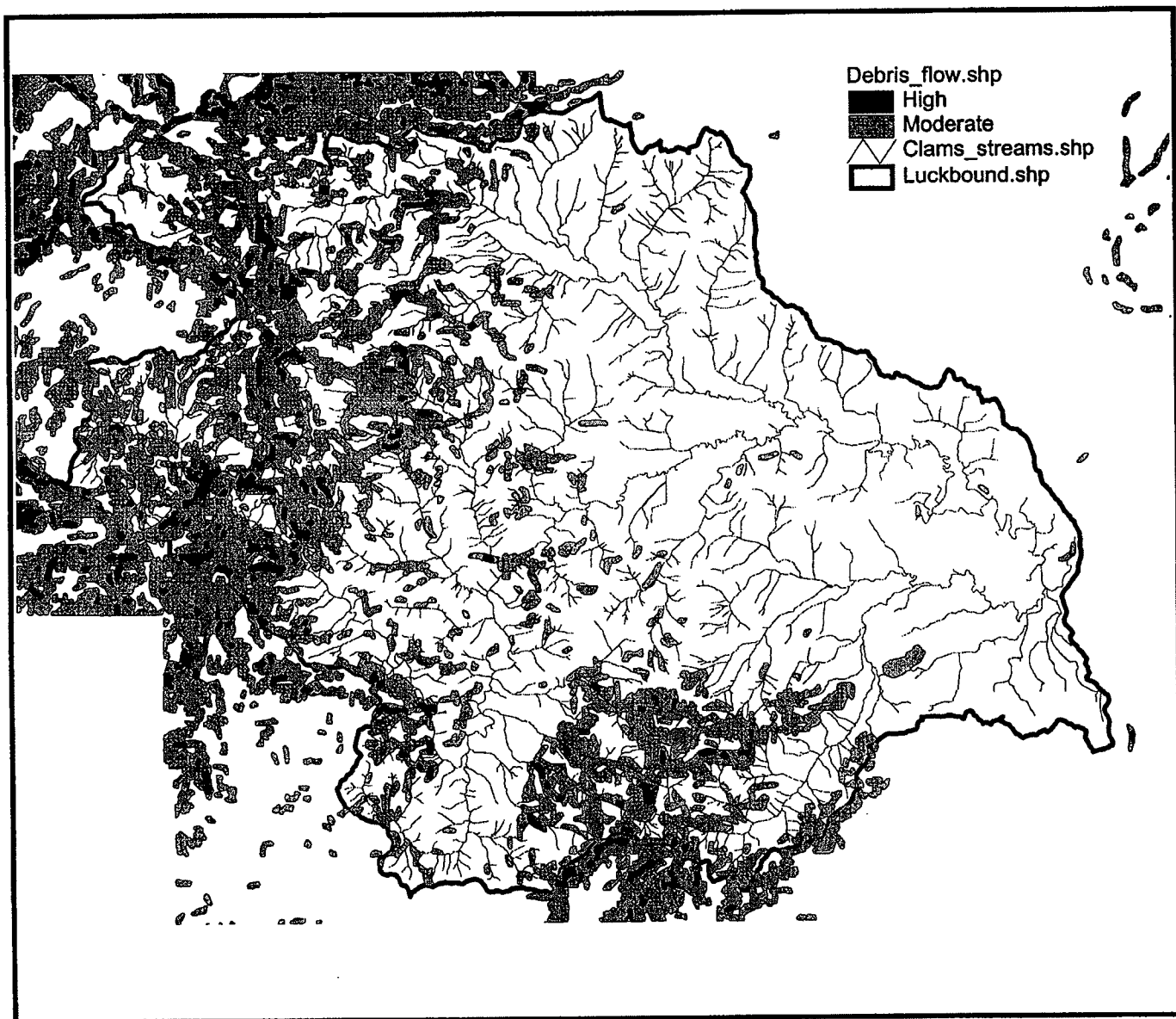
§No direct evidence for uplift; probably should be subsumed under the Bethel geomorphic surface.

**Recognized by Gelderman and Parsons (1972) near Amity (Fig. 14).

‡Recognized by Glasmann and others (1980) in a small study site 3 miles south of Monmouth (Fig. 14), but probably widespread.

§§Green (1983), Parsons (1985).

Luckiamute Basin Debris Flow Hazard Zones (OR Forestry)



5000 0 5000 10000 Meters

Table 2. Four-fold surficial map protocol for unglaciated mountainous landscapes (after Taylor and others, 1996).

1. Age of surficial material

H = Holocene (< 10,000 years old)
 W = Wisconsin (89 to 10 ka)
 I = Illinoian
 P = Pleistocene undifferentiated
 EP = early Pleistocene
 MPI = middle Pleistocene
 LP = late Pleistocene
 Q = Quaternary undifferentiated
 CZ = Cenozoic undifferentiated

2. Origin of surficial process

A. Hillslope
 r = residuum (in situ regolith)
 c = colluvium (mass wasting)
 ds = debris slide
 rf = rock fall or topple

B. Valley bottom
 a = stream alluvium (normal flow)
 hcf = hyperconcentrated flow
 df = debris flow
 sw = slackwater deposition

C. Lacustrine
 l = lacustrine deposit, undiff.
 lb = lake-bottom deposit
 ld = lacustrine deltaic

D. Other
 g = glaciofluvial, undifferentiated
 go = glacial outwash
 e = eolian
 cr = cryoturbation
 x = anthropogenic disturbance
 f = artificial fill
 rk = bedrock

3. Landform Units

A. Hillslope
 n = nose
 sl = side slope
 h = hollow
 veneer = < 2 m of regolith
 blanket = > 2 m of regolith
 bf = boulder field
 bs = boulder stream
 pg = patterned ground
 tfs = talus

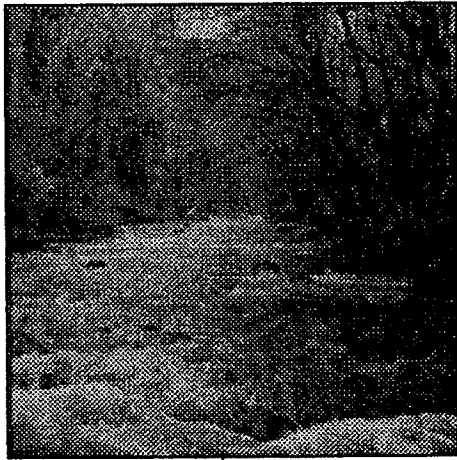
B. Valley bottom
 ch = channel
 fp = floodplain (R.I. $\leq 2-3\text{ yr}$)
 t = terrace (t1, t2 ...tn; height above river)
 f = fan
 f-t = fan terrace (f1, f2 ...fn; height above river)
 a = apron (footslope deposit)
 lo = lobe
 lv = levee
 ox = oxbow, abandoned channel

C. Other
 ft = flow track (debris flow)
 hm = hummocky topography
 rb = rock block-slide deposit
 x = excavated, fill, disturbed ground
 d = delta

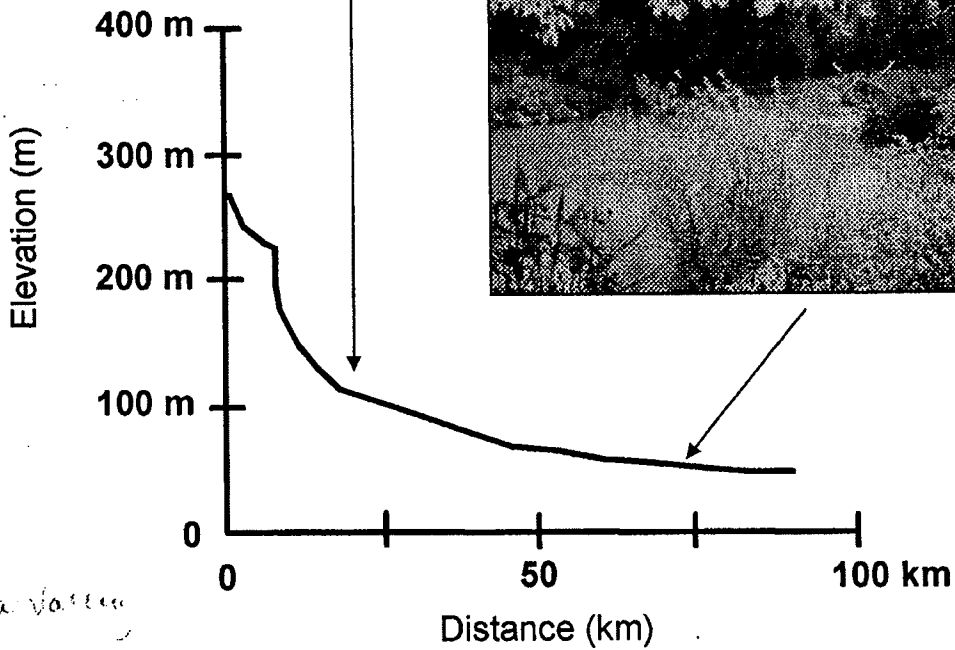
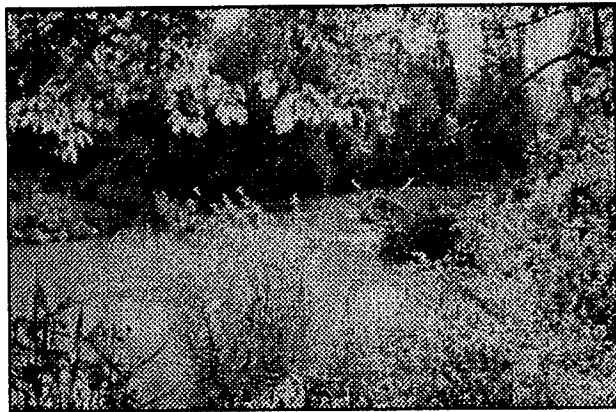
4. Material (composition and texture)

b = boulders (>256 mm clast supported)
 c = cobbles (64-256 mm clast supported)
 p = pebbles (4-64 mm clast supported)
 g = gravel (>2 mm clast supported)
 sg = mixed sand and gravel
 s = sand (0.05-2.0 mm)
 st = silt (0.002-0.05 mm)
 cy = clay (<0.002 mm)
 l = loam (mix of sand, silt, clay)
 d = diamicton undifferentiated
 bbd = very bouldery diamicton
 bd = bouldery diamicton
 cd = cobbly diamicton
 pd = pebbly diamicton
 ds = sandy matrix diamicton
 dt = silty matrix diamicton
 dy = clayey-matrix diamicton
 rk = bedrock (modify by lithology)
 rs = rotten stone, saprolite
 tr = travertine
 tu = tufa
 ma = marl
 og = organic-rich sediment
 w = water
 u = unknown

Bedload Channel



Suspended-Load Channel



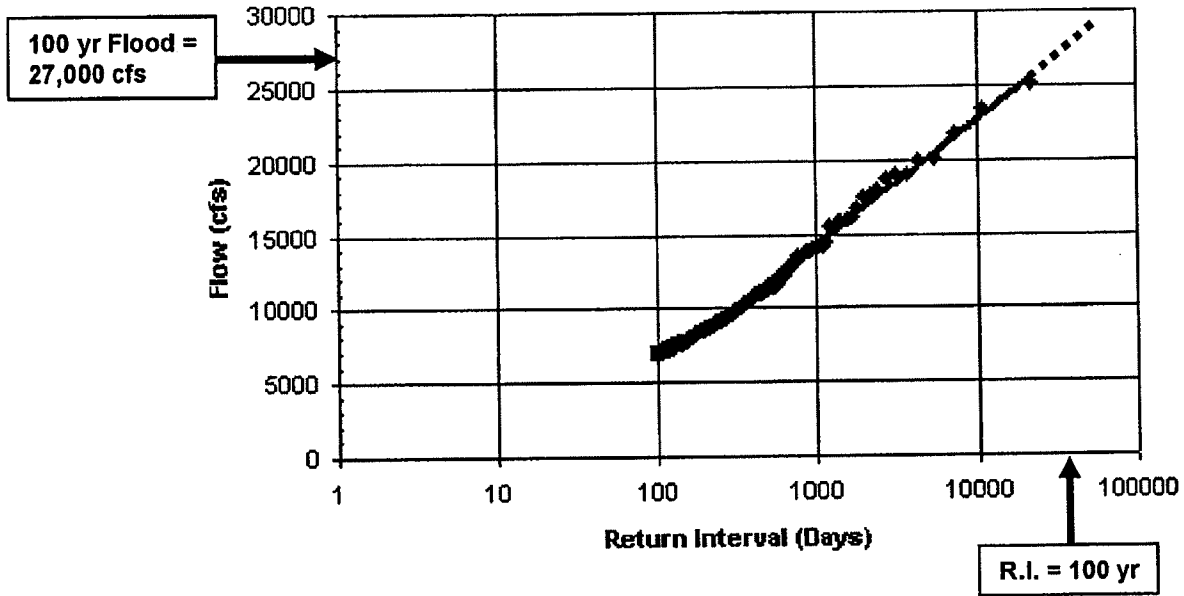
Let's look at the longitudinal profile of the Luckiamute River.

Lilliamute Valley

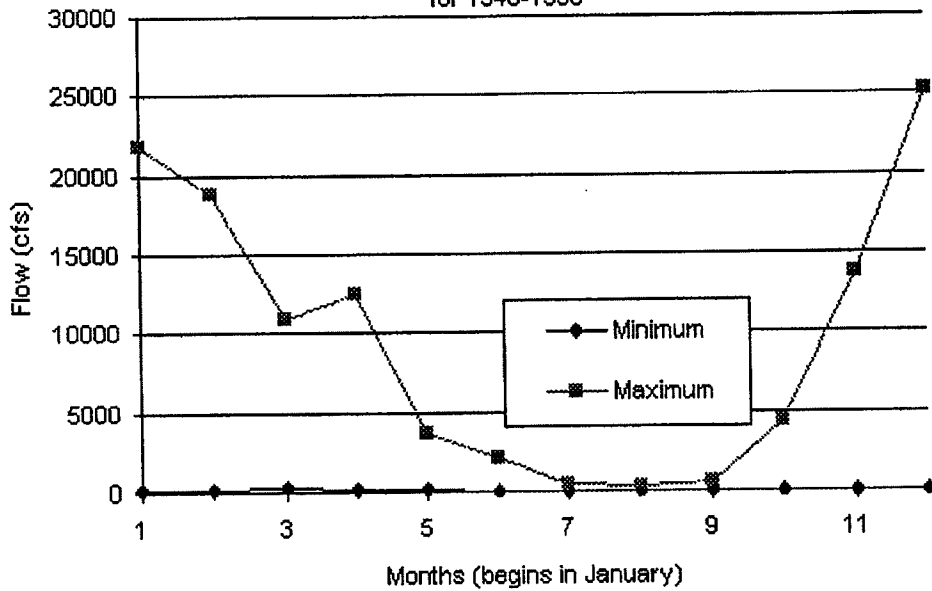
Longitudinal profile along the Luckiamute River (from Rhea, 1993).
Photos from Waichler and others (1997).

21

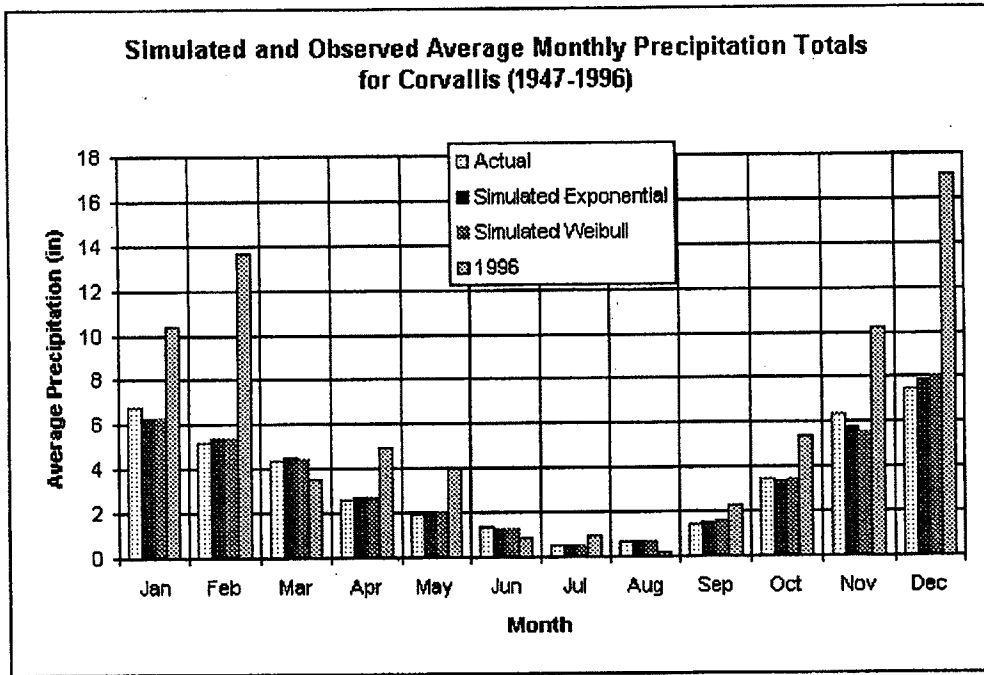
**Streamflow Return Intervals
Luckiamute R. at Suver**



**Luckiamute R. Flow at Suver
Minimum and Maximum by Month
for 1940-1996**



**Discharge characteristics for Luckiamute River, Helmick State Park
(from Waichler and others, 1997).**



(Waichler and others, 1997)

in. Waichler and others indicate precipitation is 60% of the average

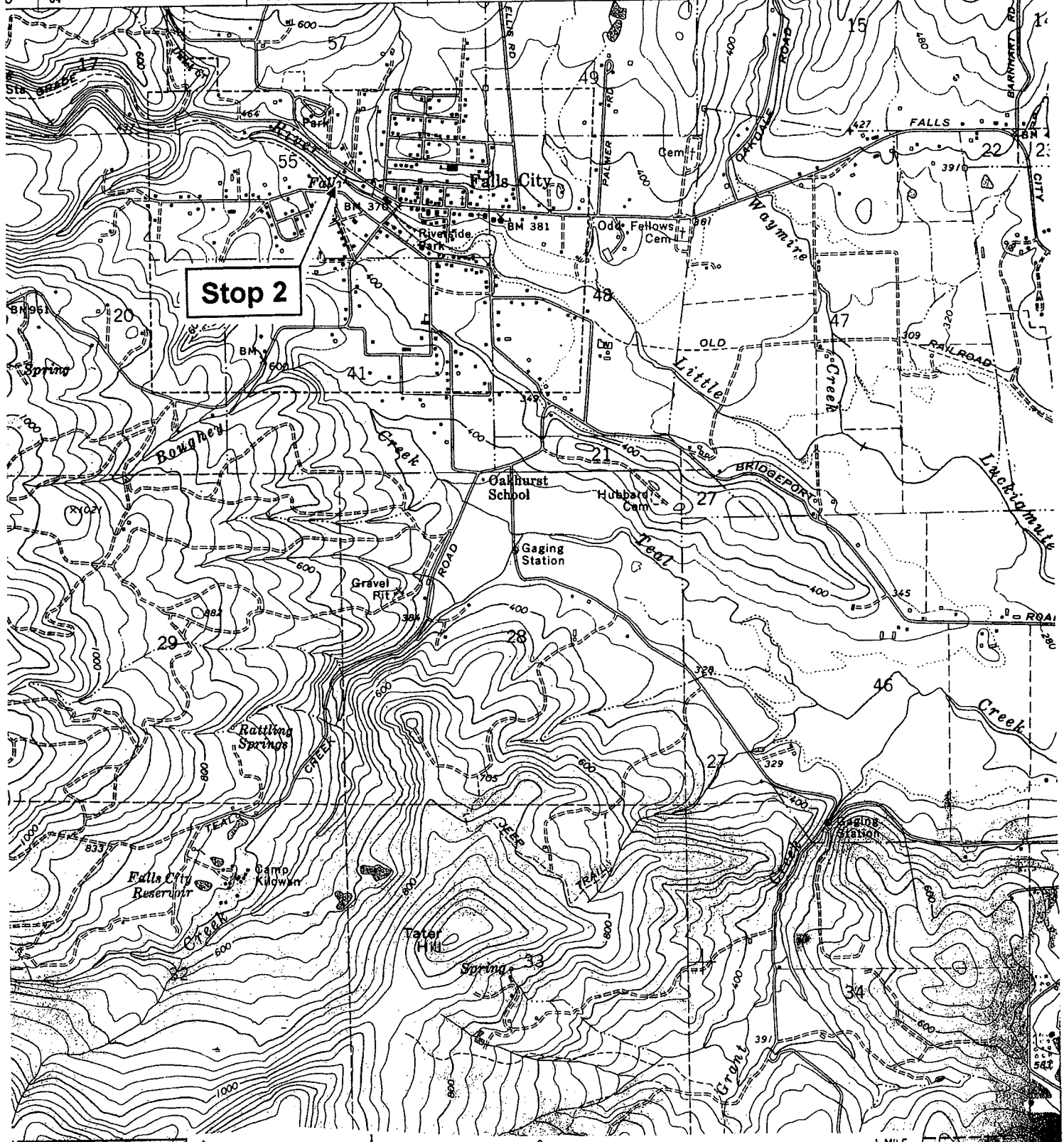
Water Balance of Luckiamute Watershed

Period (1961-1990)	Mean Precipitation (mm)	Precip. (Input) (m ³)	Observed Mean Discharge (cfs)	Observed Total Discharge (m ³)	Difference (Precip-Discharge) (m ³)	Difference as % of Precip. ("%LOSS")
Annual	1894	1.23E+09		7.55E+08	4.77E+08	39%
Jan	335	2.18E+08	2232.146	1.69E+08	4.86E+07	22%
Feb	258	1.68E+08	1853.278	1.27E+08	4.09E+07	24%
Mar	216	1.41E+08	1472.097	1.12E+08	2.89E+07	21%
Apr	101	6.57E+07	795.9956	5.84E+07	7.27E+06	11%
May	51.9	3.38E+07	396.072	3.00E+07	3.72E+06	11%
Jun	41.7	2.71E+07	188.61	1.38E+07	1.33E+07	49%
Jul	11.3	7.35E+06	71.32473	5.41E+06	1.94E+06	26%
Aug	23.8	1.55E+07	37.25441	2.83E+06	1.27E+07	82%
Sep	50.3	3.27E+07	49.19311	3.61E+06	2.91E+07	89%
Oct	143	9.30E+07	124.4226	9.44E+06	8.36E+07	90%
Nov	284	1.85E+08	904.1411	6.64E+07	1.18E+08	64%
Dec	378	2.46E+08	2069.228	1.57E+08	8.89E+07	36%

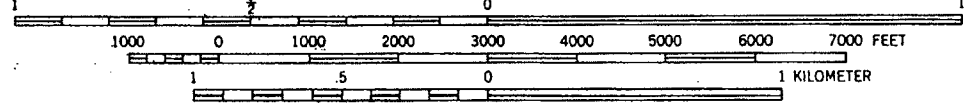
Water budget for Luckiamute Watershed (Waichler and others, 1997)

Stop 2. Falls at Falls City

- Physiographic Location
 - Little Luckiamute River (northern watershed)
- Bedrock and Surficial Geology
 - Stop is located at prominent knick point of Little Luckiamute
 - Stop is located in the Yamhill-Intrusive lithospatial domain (Figure 3, p. 10)
- Content Piece – Fluvial Erosion Dynamics
 - Knickpoint = hydraulic step in gradient
 - Falls are fracture and bedrock controlled
 - Headward erosion, block plucking, and wall-rock undercutting
 - Stream Power = (Discharge) x (Gradient) x (Specific Wt.)
 - Stream Power > Load = Erosion
 - Stream Power < Load = Deposition
 - Little Luckiamute Channel Condition at Stop 2
 - Channel under capacity with respect to sediment load
 - Power > sediment load



Stop 2

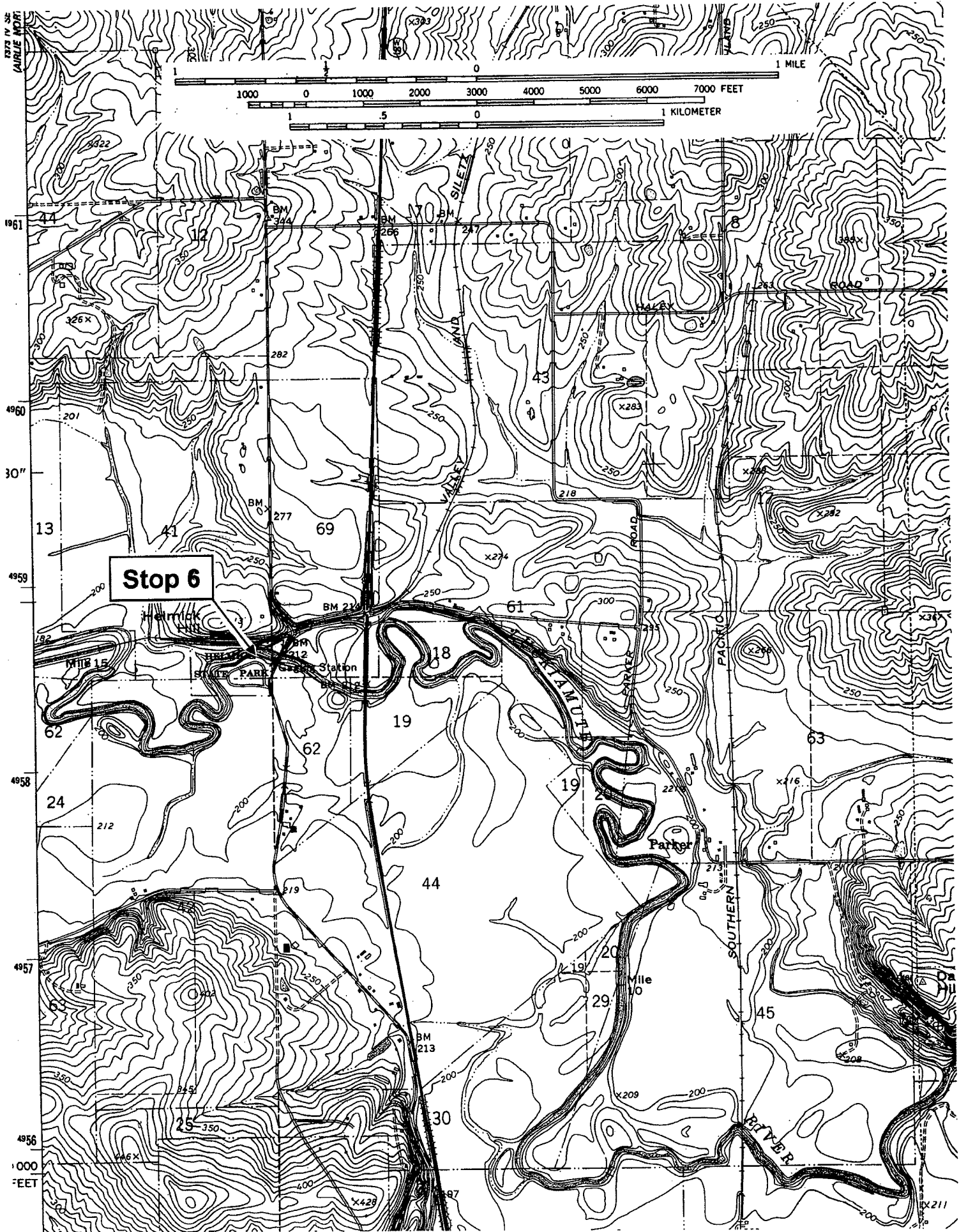


En Route to Stop 6

- Drive from Falls City, out of Yamhill-Intrusive Domain, into Spencer Fm-Valley Fill Domain (refer to Figure 3, p. 10)
- Note change in land use from forestry to agriculture
- Stop at Western Oregon University, Natural Sciences Building, for brief slide show discussion of the Polk County Flora Project and Interactive Flora Identification Key

Stop 6. Helmick State Park

- Physiographic Location
 - Lower Luckiamute River, 18 km upstream from watershed outlet into the Willamette River
- Bedrock and Surficial Geology
 - Stop is located in the Spencer Fm-Valley Fill lithospacial domain (Figure 3, p. 10)
 - Note incised channel characteristics and low terraces
- Content Piece – Field Botany and Flood Hydrology
 - Field Botany
 - Participants are provided an opportunity to use Dutton's interactive flora identification key
 - Flood Hydrology
 - USGS Suver Gaging Station
 - Recurrence intervals and seasonal discharge patterns (p. 32)
 - Paleoflood hydrology



Stop 6